

PALEOENVIRONMENTAL ANALYSES OF THE KIRKWOOD FORMATION

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ABSTRACT

Palynological analyses were made on 200 samples of the Kirkwood Formation from 7 wells and 3 outcrop localities to determine the environments of deposition and climatic conditions that prevailed during its deposition.

The Kirkwood Formation consists of three surface members: the Asbury Park, Grenloch Sand and Alloway Clay. Five subsurface phases, identified and described in this study are correlated with the Early Miocene Tampa and Hawthorn Formations of South Carolina, Georgia, and Florida, and the Middle Miocene Calvert, Choptank and St. Marys Formations of Maryland, Delaware and Virginia. The Asbury Park, Grenloch Sand and Alloway Clay Members are correlated with the subsurface Calvert Phase on the basis of their palynological assemblages.

Paleobathymetric interpretations of the surface and subsurface facies of the Kirkwood Formation are made on the basis of their pollen, spore, dinoflagellate, diatom, silicoflagellate and radiolarian assemblages. These assemblages indicate nearshore environments of deposition for the Asbury Park, Grenloch Sand and Alloway Clay Members and transgressive and regressive sequences in the Tampa, Hawthorn, Calvert and St. Marys Phases.

Paleoclimatic interpretations, made on the basis of Pinus/Picea ratios and changes in the regional microfloras indicate that the Asbury Park, Grenloch Sand and Alloway Clay

were deposited during a period of climatic deterioration from subtropical to temperate conditions. A period of climatic deterioration from subtropical to temperate conditions is indicated to have taken place during the deposition of the Tampa, Hawthorn and lower portion of the Calvert Phase. A period of climatic amelioration is indicated during the deposition of the upper part of the Calvert Phase and throughout the deposition of the Choptank Phase. A second period of climatic deterioration is indicated to have taken place during the deposition of the St. Marys Phase.

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PREVIOUS PALEOENVIRONMENTAL STUDIES

Paleoenvironmental interpretations of the surface samples of the Kirkwood Formation have been made by several authors on the basis of paleontological, sedimentological and mineralogical studies.

Isphording and Lodding (1970 and 1973) determined the kaolinite, illite and montmorillonite content of the Asbury Park Member and concluded that it was deposited in transitional marine environments such as marshes, lagoons or estuaries.

Richards and Harbison (1942) studied the molluscan faunas of the Shiloh Marl facies of the Alloway Clay Member and placed it in the middle neritic environment of deposition.

Isphording and Lodding (1973) reported that the Macro-kaolinite Zone of the Alloway Clay was deposited in either a lagoon, swamp or estuary that received large amounts of runoff.

Paleoclimatic interpretations of the Kirkwood Formation have been reported by Isphording (1970) and Isphording and Lodding (1973). They reported that the surface members of the Kirkwood Formation were deposited under warm, moist climatic conditions. Their interpretations were based on heavy mineral, light mineral and clay mineral suites.

Prior to the present study, paleobathymetric and paleoclimatic interpretations had not been reported on the alternating sands and clays of the Kirkwood Formation in the subsurface.

Paleoenvironmental interpretations of the Kirkwood Formation in the present study are made on the basis of microfossils recovered from 200 samples from 7 wells and 3 outcrop localities. Included in these assemblages are pollen, spores, silicoflagellates, radiolaria, dinoflagellate cysts, diatoms and foraminiferal test linings.

Paleobathymetric interpretations are based on spore/dinoflagellate ratios as well as relative frequency distributions of continental and marine palynomorphs. In general, spore/dinoflagellate ratios decrease from transitional to nearshore and offshore deposits (Dunay 1969).

Previous palynological studies of other Tertiary formations of North America are limited in that their assemblages are often exclusively marine or continental. Continental assemblages are limited in that they provide short term discontinuous records of the vegetative history of the area. These assemblages often represent the flora of a limited area in the immediate area of the deposit (Groot and Groot 1966). Thus it is difficult to ascertain long term climatic fluctuations from these assemblages.

An additional problem has been the correlation of previously described, isolated palynological assemblages. The latitudinal zonation of the tundra, boreal, mixed deciduous and deciduous plant communities across North America is well documented (Potszger and Otto 1943, Goodlet 1954, Graham 1963, Davis 1969 and Elsik 1969). Because of this latitudinal zonation it is virtually impossible to correlate

the continental Miocene deposits of British Columbia as reported by Piel (1969) for example with those of Louisiana as reported by Elsik (1969) on the basis of palynological assemblages alone.

Exclusively marine palynological assemblages are limited in that the asaccate paleoclimatical indicators that appear in abundance in continental assemblages are often rare to absent in marine assemblages (Dunay 1969).

Because of the absence of asaccate grains in offshore assemblages, paleoclimatic interpretations should be made on the basis of the relative frequencies of arboreal forms that produce bisaccate grains and whose distributions in modern floras are controlled by climatological factors. Paleoclimatic interpretations in this study are based on Pinus/Picea ratios as well as overall changes in the regional floras.

Modern species of Pinus are included in the floras of many regions throughout North America (Munn 1938, Braun 1950, Fernald 1950 and Collingwood and Brush 1964). Associated forms, including Taxus and Taxodium indicate that the species of Pinus that are present in palynological assemblages of the Kirkwood Formation probably had climatic tolerances and ecologic requirements similar to modern species of Pinus that are now included in the native floras along the Gulf Coastal Plain and the southern portion of the Atlantic Coastal Plain. These forms include Pinus echinata Miller, Pinus palustris Miller, Pinus ellioti Engelm, Pinus taeda Linneaus, Pinus

serotina Michaux, Pinus virginia Miller and Pinus glabra Walter (Fernald 1950 and Collingwood and Brush 1964).

Modern species of Picea closely follow the 10°C average July isotherm on their northern boundaries and the 23°C isotherm on their southern boundaries (Wolfe and Leopold 1967). Picea rubens Sargent and Picea mariana (Miller) are major elements of the modern floras from northern New England to the Yukon territory (Collingwood and Brush 1964). Pollen grains of these species very closely resemble the Picea grains that are encountered in palynological assemblages of the Kirkwood Formation. Other occurrences of Picea in the modern floras of North America include the appearances of Picea sitchensis (Bongard) Carriere and Picea breweriana Watson in the modern flora of the Coastal Ranges from northern California to Alaska and Picea pungens Engelmann and Picea engelmanni Perry on the higher slopes of the Northern Rockies (Munn 1938, Braun 1950, Fernald 1950 and Collingwood and Brush 1964).

Increasing Pinus/Picea ratios in Quaternary floras indicate warming conditions. Decreasing Pinus/Picea ratios indicate cooling conditions (Davis 1967).

The Kirkwood Formation consists of four alternating marine and terrestrial facies in the subsurface. Widely spaced sections are correlated with each other on the basis of their palynological assemblages. Thus a continuous regional paleoenvironmental analysis may be made from samples of the Kirkwood Formation.

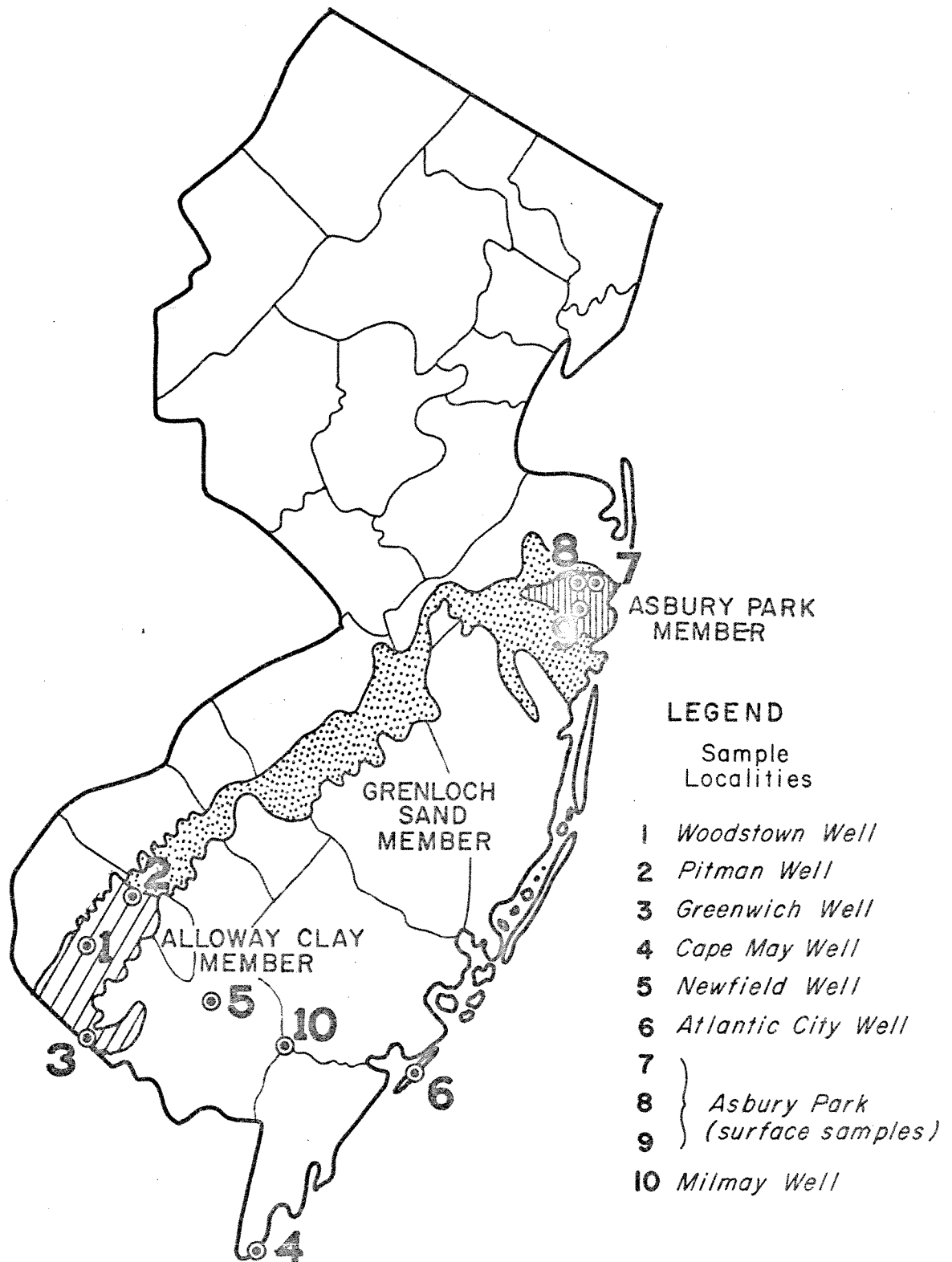


Figure 1: Outcrop map of the Kirkwood Formation including sample locations of the present study

STRATIGRAPHY OF THE KIRKWOOD FORMATION

The Kirkwood Formation outcrops in a northeast to southwest trending belt on the New Jersey Coastal Plain. The formation dips 10 to 25 feet per mile below the overlying Cohansey Formation (Barksdale 1958). The Kirkwood Formation unconformably overlies the Shark River and Manasquan Formations (Eocene) (Richards 1945 and Owens and Sohl 1973). The maximum outcrop thickness is 100 feet. The formation thickens to 790 feet in the subsurface at Atlantic City (Barksdale 1958).

The Kirkwood consists of three members, the Asbury Park, Grenloch Sand and Alloway Clay.

The Asbury Park Member outcrops in Monmouth County (Fig. 1). It consists of dark brown, often finely laminated, silty sands and micaceous clays.

The Alloway Clay outcrops in Cumberland, Gloucester and Salem Counties (Fig. 1). The clay is light brown in outcrop and dark brown to gray in the subsurface. The upper beds of the Alloway Clay near Shiloh, New Jersey are locally called the Shiloh Marl. In the vicinity of Woodstown, New Jersey, the illite and montmorillonite near the base of the Alloway Clay have been diagenetically altered to form a "macro-kaolinite zone" (Isphording and Lodding 1970 and 1973). This unit seems to be unique in the geological literature. It had for many years been mistakenly referred to as "micaceous talc-like clay" (Ries and Kummel 1904).

The Grenloch Sand is the largest of the three members in areal extent. It outcrops in portions of Monmouth, Ocean, Burlington, Camden, Gloucester and Salem Counties (Fig. 1). This member overlies and is interlayered with the Asbury Park Member to the northeast and the Alloway Clay Member to the southwest (Fig. 2). In outcrop the Grenloch Sand Member consists of fine yellow and orange sands. In the subsurface the sands are interlayered with silts and clays.

The Asbury Park, Grenloch Sand and Alloway Clay are easily recognizable in outcrop, but rapidly lose their identities downdip and cannot be differentiated lithologically in deep wells (Isphording and Lodding 1970).

In the present study subsurface sections of the Kirkwood Formation are divided into five biostratigraphic units. Units I and II are encountered in deep well sections from Greenwich, Newfield, Milmay and Cape May (Figs. 12-16). Unit I consists of gray silty clays and varies in thickness from 50 feet at Greenwich to 70 feet at Cape May. Unit II consists of gray to brown silty clays and varies in thickness from 25 feet at Greenwich to 60 feet at Cape May. Units III, IV and V are encountered in sections from Greenwich, Newfield, Milmay, Atlantic City and Cape May (Figs. 12-16). Unit III consists of alternating sands and clays and varies in thickness from 45 feet at Greenwich to 220 feet at Cape May. Unit IV consists of sands and clays and varies in thickness from 20 feet at Greenwich to 150 feet at Cape May. Unit V

consists of alternating sands and silts and varies in thickness from 40 feet at Greenwich to 200 feet at Cape May.

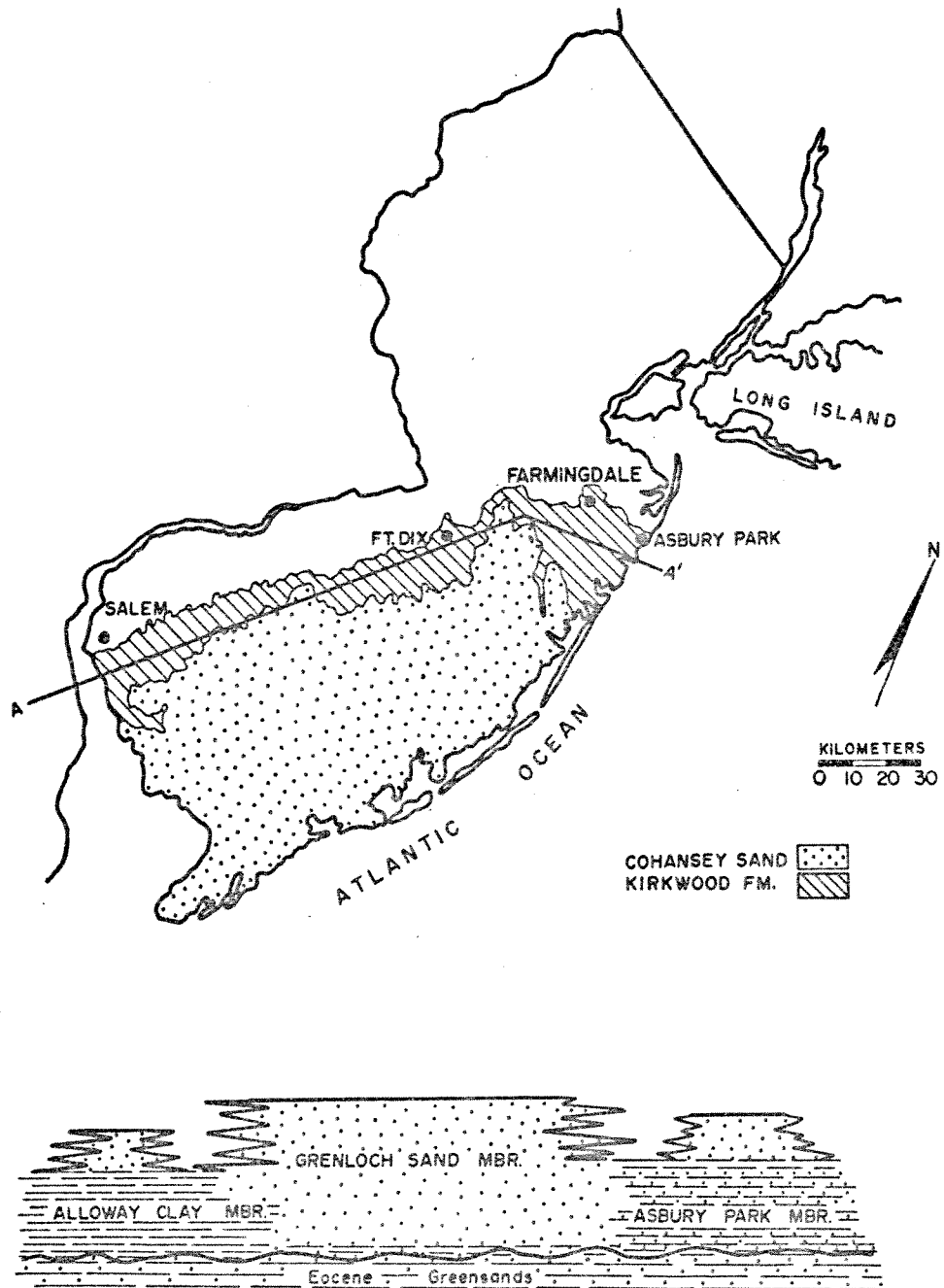


Figure 2: Facies relationships of the Asbury Park, Grenloch Sand and Alloway Clay Members of the Kirkwood Formation

RESULTS OF PALYNOLOGICAL ANALYSES OF THE SURFACE MEMBERS OF
THE KIRKWOOD FORMATION

Palynological assemblages of the Asbury Park Member are dominated by nonarboreal palynomorphs including those of Chenopodiaceae, Ericaceae and Gramineae. Spores of lower vascular plants including Equisetum are abundant in most samples. Arboreal associates include Alnus, Betula, Castanea, Corylus, Cyrilla, Fagus, Juniper, Liquidambar, Picea, Pinus, Quercus, and Taxodium. Aquatic elements include dinoflagellates, diatoms and silicoflagellates. The diatom, Actinoptychus heliopelta Grunow and the dinoflagellate, Deflandrea are present in samples collected from Farmingdale, Hammonton and Howell.

Surface samples of the Grenloch Sand are void of pollen, spores, silicoflagellates, diatoms, radiolarians, dinoflagellate cysts and foraminiferal test linings. Palynological residues of the Grenloch Sand consist of highly oxidized fragments of plant tissues.

Palynological assemblages of the Alloway Clay Member are dominated by arboreal pollen grains. Pinus, Picea, Quercus, Carya, Juniper, Fagus, Acer and Taxodium are the dominant forms. Nonarboreal palynomorphs including members of Chenopodiaceae, Ericaceae and Gramineae are present in the upper portion of this member. Aquatic elements consisting primarily of large dinoflagellate cysts, including Hystriochokolpoma and foraminiferal test linings become less

abundant in the upper part of the section. Palynological assemblages of the "macrokaolinite zone" consist almost exclusively of Pinus grains. Marine, nonarboreal and arboreal forms including Picea, Quercus, Carya and Tilia are rare in these assemblages.

RESULTS OF PALYNOLOGICAL ANALYSES OF THE SUBSURFACE PHASES OF
THE KIRKWOOD FORMATION

Subsurface sections of the Kirkwood Formation may be divided into five biostratigraphic units on the basis of their palynological assemblages. Units I and II are encountered in deep well sections from Greenwich, Newfield, Milmay and Cape May (Fig. 1). Units III, IV and V are encountered in sections from Greenwich, Newfield, Milmay, Atlantic City and Cape May (Fig. 1).

Palynological assemblages in the lower portion of Unit I are dominated by large dinoflagellate cysts including those of Hystriochokolpoma. Pinus, Quercus and Carya are the dominant arboreal forms throughout these assemblages. Alnus, Betula, Castanea, Corylus, Cyrilla, Ilex, Nyssa, Platycarya, Pterocarya, Salix, Taxodium and Taxus are minor arboreal constituents. Frequencies of nonarboreal elements including members of Chenopodiaceae, Ericaceae and Gramineae generally increase in the upper part of the section.

Palynological assemblages of the lower portion of Unit II are dominated by small spiny dinoflagellate cysts including forms closely resembling modern species of Micrhystriidium. Arboreal and nonarboreal elements are similar in composition and distribution to those of Unit I.

Pinus, Quercus, Carya and dinoflagellate cysts dominate the palynological assemblages of Unit III. Frequencies of dinoflagellate cysts decrease in the upper part of the section

with small spiny forms becoming relatively more abundant. The first major influxes of Picea and Gramineae occur in this unit. Minor arboreal elements include Alnus, Betula, Castanea, Corylus, Cyrilla, Ephedra, Fagus, Ilex, Liquidambar, Magnolia, Melia, Nyssa, Platanus, Platycarya, Salix, Taxodium, Taxus and Tilia. Nonarboreal forms including high frequencies of Gramineae and unidentified spores of lower vascular plants increase in the upper portion of Unit III. Foraminiferal test linings, radiolarians, diatoms and silicoflagellates including Corbisema and Dichtyocha fibula are abundant in the lower portion of this unit.

Palynological assemblages of Unit IV are dominated by Quercus, Pinus and Carya. Other arboreal forms include Alnus, Betula, Castanea, Corylus, Cyrilla, Engelhardtia, Fagus, Juglans, Nyssa, Salix, Taxodium, Tilia and Ulmus. Relative frequencies of Picea decrease in the upper portion of this unit. Picea completely disappears in the upper portion of this unit from assemblages taken from Greenwich, Milmay and Atlantic City. Nonarboreal elements including Chenopodiaceae, Ericaceae, Gramineae and spores of lower vascular plants are abundant in the central portion of this unit. Marine forms consisting primarily of small, spiny dinoflagellate cysts first decrease and then increase in the upper portion of the section.

Palynological assemblages of Unit V are dominated by Pinus, Quercus and Carya. Minor arboreal constituents include Acer, Alnus, Betula, Castanea, Corylus, Cyrilla, Fagus,

Juglans, Liquidambar, Magnolia, Nyssa, Pterocarya, Prunus, Taxodium, Tilia and Ulmus. Picea appears in low frequencies in lower portions of this unit and increases in the upper levels. The first appearance of Abies is noted in this unit in samples taken from Greenwich, Newfield, Atlantic City and Cape May. Nonarboreal elements including members of Chenopodiaceae, Ericaceae, Gramineae, spores of lower vascular plants, leaf hairs and Sporangium are abundant in the upper portion of this unit. The first appearance of Compositae is noted in assemblages from Greenwich, Milmay, Atlantic City and Cape May. Marine elements consisting primarily of dinoflagellate cysts generally decrease with smaller forms becoming relatively more abundant than larger forms in the upper portion of the unit.

Percentage frequency histograms of the palynological analyses are presented in Appendix II.

PALEOENVIRONMENTAL INTERPRETATIONS OF THE SURFACE MEMBERS OF
THE KIRKWOOD FORMATION

Palynological assemblages of the Asbury Park Member include high frequencies of Chenopodiaceae and Ericaceae. Modern species of Chenopodiaceae are found in sandy waste soils. The phreatic waters of these soils are often saline or brackish (Fernald 1950). Modern species of Ericaceae are often found in peaty clearings and along the margins of bogs (Fernald 1950).

The dinoflagellate, Deflandrea, a common form in modern nearshore brackish water assemblages, is included in several palynological assemblages of the Asbury Park Member.

The association of members of Chenopodiaceae and Ericaceae with dinoflagellate cysts indicates a coastal brackish water environment of deposition similar to that reported by Martin and Rouse (1966) from the Oligocene deposits of the Queen Charlotte Islands of British Columbia.

Pinus/Picea ratios ranging from 7/1 to 10/1 and the presence of Taxodium in palynological assemblages of the Asbury Park Member indicate that these sediments were deposited under subtropical climatic conditions. Modern species of Taxodium are found in the native floras of the Gulf Coastal Plain and the southern portion of the Atlantic Coastal Plain (Fairchild and Elsik 1968, Fredrickson 1969, Stewart 1971 and Tschudy 1973).

Surface samples of the Grenloch Sand are void of micro-paleontologic remains with the exception of highly oxidized fragments of maceral tissues. The association of these oxidized macerals along with the relatively coarse texture of these deposits indicates a high energy, possibly littoral, environment of deposition for the Grenloch Sand Member.

Palynological assemblages of the Alloway Clay Member are characterized by decreasing frequencies of large dinoflagellate cysts and the presence of nonariboreal forms in the upper portion of the unit. These assemblages, including high frequencies of planktonic foraminiferal test linings in the lower portion of the unit, indicate a middle neritic environment of deposition becoming more shallow, possibly to inner neritic, during the final stages of its deposition (Wall 1965, Groot and Groot 1966 and Davey 1970a).

Decreasing Pinus/Picea ratios from 9/1 to 6/1 and the presence of modern subtropical forms including Taxodium in the lower portion of the Alloway Clay indicate that climate deteriorated during the deposition of this unit. It appears that the subtropical conditions that prevailed during the deposition of the lower portion of this member gave way to temperate conditions during the deposition of its upper portion, the Shiloh Marl. Increasing Pinus/Picea ratios indicate a warming trend during the final stages of the deposition of the Shiloh Marl. The presence of Taxus, limited in modern assemblages to the southern portion of Florida, in the uppermost horizon of the Shiloh Marl at

Pitman, New Jersey supports the warming trend interpretation suggested by the Pinus/Picea ratios.

The apparent lack of arboreal grains other than Pinus in palynological assemblages of the Macrokaolinite Zone of the Alloway Clay has been explained by Isphording and Lodding (1973) to be the result of the diagenesis that produced the macrokaolinite. They reasoned that the other pollen grains that were present in the Macrokaolinite Zone were dissolved during the diagenesis while the hardier Pinus varieties were unaffected. This seems not to be the case because, although rare, well preserved grains produced by Carya, Picea, Quercus and Tilia are present in these assemblages.

It is more probable that the predominance of Pinus is due to its anemophilous form of dispersal and the relatively large quantities of pollen that each pine tree produces. In modern palynological assemblages in the general vicinity of pine forests, frequencies of Pinus are over-represented from their actual numbers by a factor of from 6.6 to 1 to 22.4 to 1 (Davis, Brubaker and Beiswenger 1971).

Palynological assemblages of the Macrokaolinite Zone are similar in composition to modern assemblages of pine swamps along the southern New Jersey to Virginia coasts.

PALEOENVIRONMENTAL INTERPRETATIONS OF THE SUBSURFACE PHASES
OF THE KIRKWOOD FORMATION

Palynological assemblages of the lower portion of Unit I are dominated by large dinoflagellate cysts and foraminiferal test linings. These assemblages indicate a middle neritic environment of deposition (Wall 1965). Frequencies of nonarboreal elements generally increase in the upper part of the section indicating shoaling conditions during the deposition of the upper horizons (Davey 1970a).

Pinus/Picea ratios greater than 100/1 and the presence of Platycarya, Pterocarya, Taxodium and Taxus in palynological assemblages of Unit I indicate deteriorating subtropical conditions during the deposition of this unit. Modern species of Platycarya, Pterocarya, Taxodium and Taxus are indigenous to subtropical and tropical areas of Asia and North America (Fernald 1950 and Kapp 1969).

Palynological assemblages of Unit II are similar in their overall composition to those of Unit I. Dinoflagellate cysts dominate the lower assemblages and then decrease in the upper assemblages as they do in Unit I; however the cysts of Unit II are smaller and contain more spines than do the cysts of Unit I. Also the foraminiferal test linings are less abundant in Unit II than they are in Unit I. These assemblages indicate a shoaling inner neritic environment of deposition (Wall 1965).

Pinus/Picea ratios vary from 50/1 to greater than 100/1 in Unit II. Relative frequencies of Platycarya, Pterocarya, Taxodium and Taxus are somewhat lower in Unit II than in Unit I. Thus the deteriorating climatic conditions during the deposition of Unit I appear to have continued throughout the deposition of Unit II.

The presence of silicoflagellates, radiolarians, dinoflagellate cysts and foraminiferal test linings in samples of Unit III indicates a middle to outer neritic environment of deposition for these sediments (Cornell 1969). Increasing nonarboreal elements in the upper portion of this unit as shown by spore/dinoflagellate ratios (Fig. 4) indicate shoaling conditions during the deposition of the upper portion of this unit (Dunay 1969).

Pinus/Picea ratios (Fig. 4) indicate that the climatic deterioration that began during the deposition of Unit I culminated during the deposition of Unit III. Pinus/Picea ratios varying from 9/1 to 7/1 and the presence of Gramineae and Ephedra, whose modern species are present in the southwestern portion of the United States, indicate the presence of cool (temperate) dry conditions along the New Jersey Coastal Plain during the deposition of Unit III.

Palynological assemblages of Unit IV contain few marine palynomorphs. Those forms that are present are small spiny dinoflagellates in the uppermost and lowermost horizons. Spore/dinoflagellate ratios (Fig. 3) are highest in the

central portion of Unit IV. These assemblages indicate that the regressive conditions that began in Unit III continued into Unit IV and culminated during the deposition of the central portion of that unit. Assemblages of the upper portion of Unit IV indicate a return to nearshore conditions.

Increasing Pinus/Picea ratios from 10/1 to greater than 100/1 appear to indicate warming conditions throughout the deposition of Unit IV (Fig. 4). Picea disappears completely from some of the assemblages of the uppermost horizons. The return of Taxus along with lower frequencies of Gramineae and the absence of Ephedra apparently indicates a return to warm moist climatic conditions along the New Jersey Coastal Plain during the deposition of Unit IV.

Spore/dinoflagellate ratios (Fig. 3) indicate that the transgressive conditions that began in Unit IV culminated during the deposition of Unit V. Small spiny dinoflagellate cysts dominate the assemblages of the central portion of Unit V indicating nearshore conditions. Increasing spore/dinoflagellate ratios in the upper horizons indicate a return to shoaling conditions during the deposition of those sediments.

Decreasing Pinus/Picea ratios (Fig. 4) from 10/1 to 7/1 and the first appearance of Abies in the uppermost horizons of Unit V indicate climatic deterioration throughout the deposition of the unit. Modern species of Abies are generally restricted to areas north of 45° N. latitude in North America.

Species of Abies that are present in the Rocky and Appalachian Mountains south of 45° N. latitude are restricted to elevations above 4,000 feet (Collingwood and Brush 1964). Pollen grains produced by Abies in palynological assemblages of Unit V are most similar, morphologically, to those produced by Abies balsamea the modern distribution of which is limited to the native floras of northern New England to the Yukon territory (Collingwood and Brush 1964 and Kapp 1969).

Paleoclimatic interpretations made in this study on the basis of Pinus/Picea ratios closely match those made by Norem (1956) and Dorf (1960, 1964 and 1969) on the basis of paleobotanic remains of isolated deposits throughout North America.

Paleobathymetric interpretations made in this study on the basis of spore/dinoflagellate ratios generally match those made by Richards and Harbison (1942) and Isphording and Lodding (1973) on the basis of the molluscan faunas and clay mineralogy of the surface members of the Kirkwood formation. Prior to this study, paleobathymetric interpretations had not been made on the subsurface samples of the Kirkwood.

TRANSGRESSIONS AND REGRESSIONS DURING KIRKWOOD DEPOSITION INFERRED FROM SPORES/DINOFLAGELLATES RATIOS*

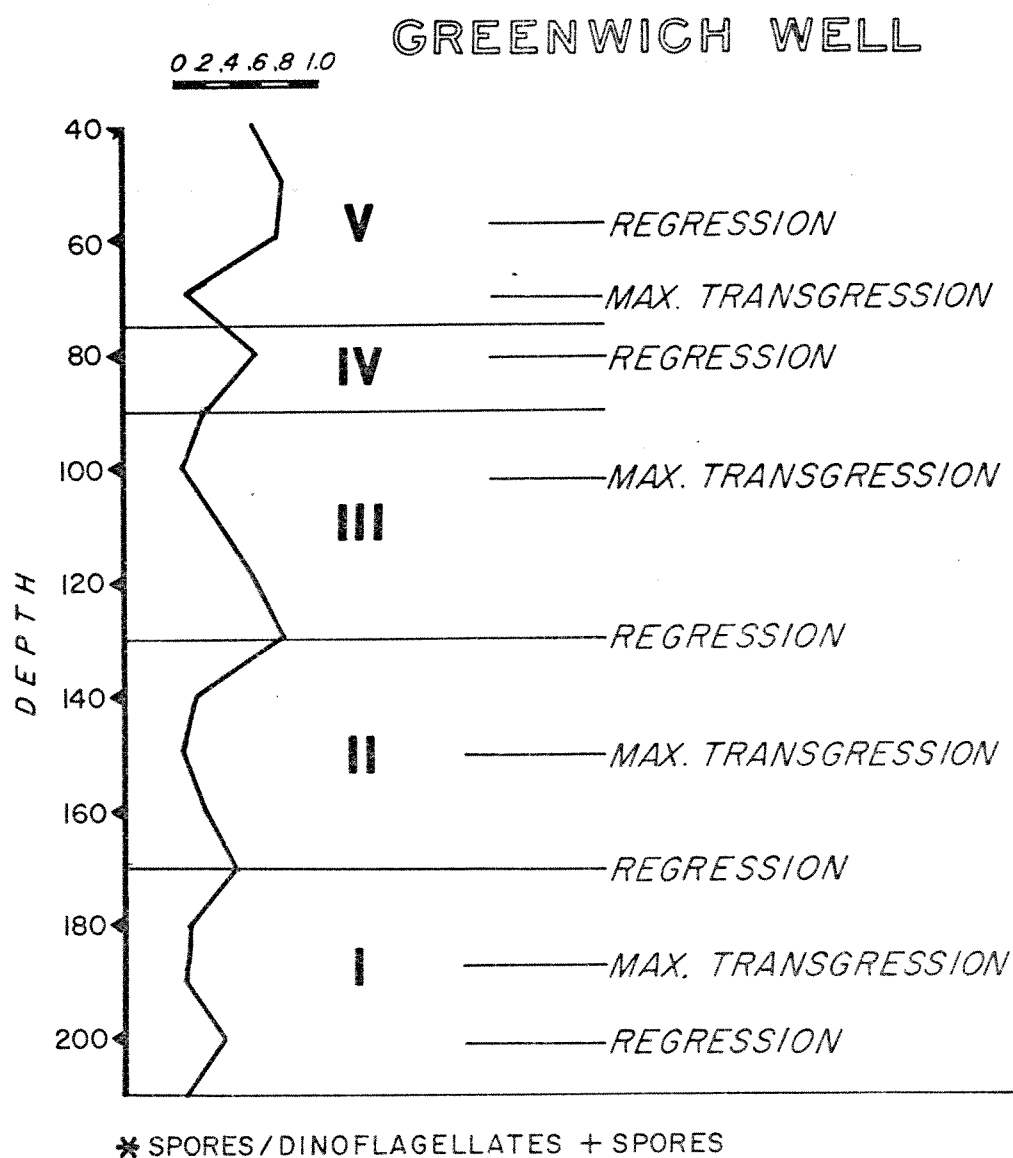


Figure 3: Spore/dinoflagellate ratios of the Kirkwood Formation at Greenwich, New Jersey

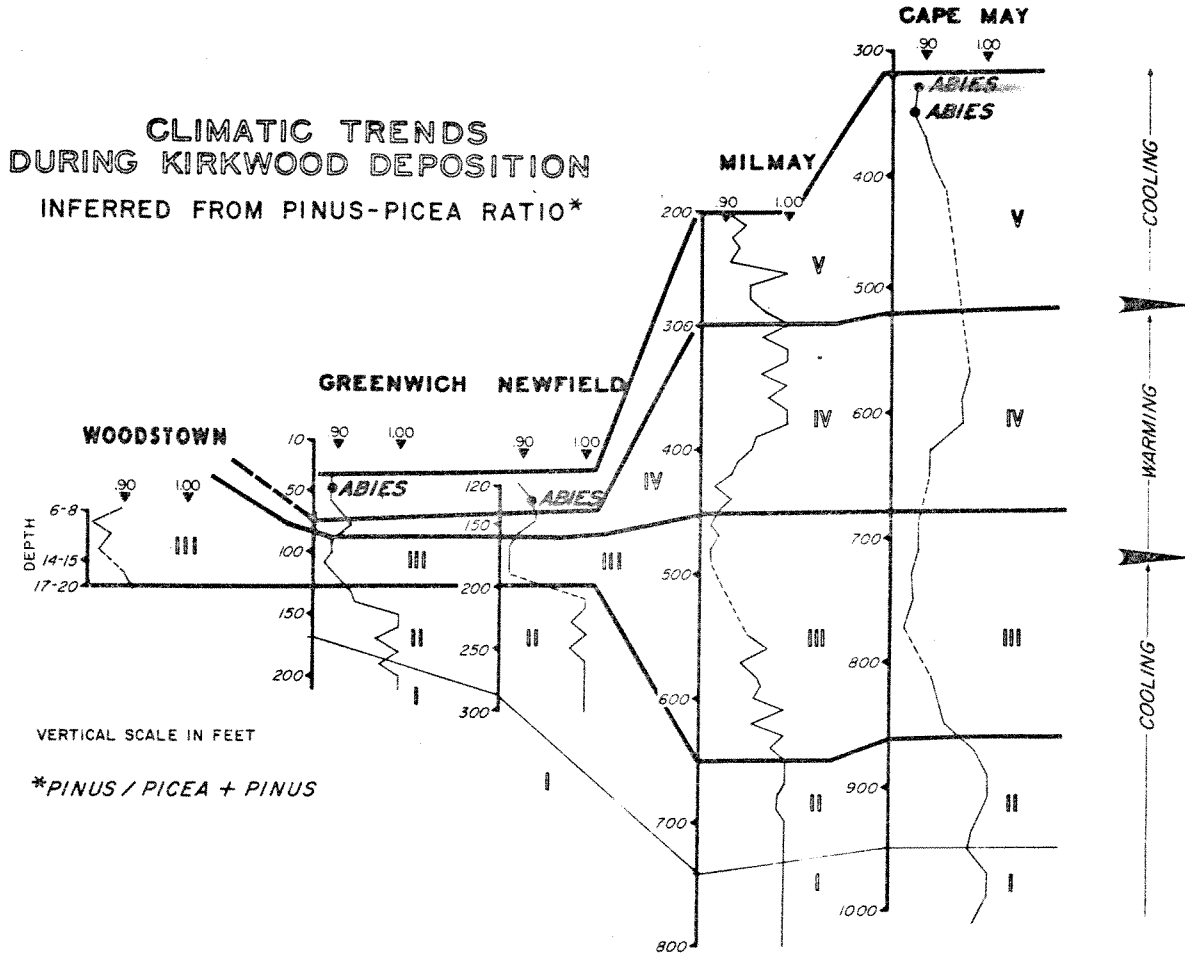


Figure 4: Pinus/Picea ratios of the Kirkwood Formation at Woodstown, Greenwich, Newfield, Milmay, and Cape May, New Jersey

CORRELATION OF THE SUBSURFACE PHASES OF THE KIRKWOOD FORMATION
WITH OTHER FORMATIONS ALONG THE ATLANTIC COASTAL PLAIN

The Miocene sediments of New Jersey were assigned to the Kirkwood Formation by Knapp (1904). Prior to this, these sediments were referred to as the Chesapeake Formation because of the similarity of its molluscan fauna to that of the Chesapeake Group of Maryland, Delaware and Virginia (Clark 1893, Whitfield 1894, and Salisbury 1895, 1896 and 1898).

Richards and Harbison (1942) correlated the Asbury Park and Alloway Clay Members with the Middle Miocene Calvert Formation, the basal member of the Chesapeake Group, on the basis of their molluscan faunas and the diatom, Actinoptychus heliopelta Grunow. Subsequent correlations by Dorsey (1948), Gardner (1948), Lohman (1948), Johnson and Richards (1952), Gernant (1970) and Owens and Sohl (1970) agree with those of Richards and Harbison (1942).

Foraminiferal assemblages of Units I and II of the present study including Cassigerinella chipolensis, Globigerina bradyi, Globigerina ouachitaensis ciperensis, Globigerina praebuloides, leroyi, Globigerina cf. quinqueloba, Globigerina venezuelana, Globigerinita dissimilis, Globigerinoides triloba and unidentified species of Globorotalia indicate that these units were deposited during the Early Miocene (Globorotalia kugleri Zone).

Unit III of the present study can be correlated with the Calvert Formation on the basis of the silicoflagellates, Corbisema sp. and Dictyocha fibula. These forms identified from the 120 foot interval in the Greenwich Well are diagnostic forms of the Calvert Formation (Tynan 1957).

Richards (1945) reported the presence of Bulliopsis integra (Conrad) and Terebra inornata Whitfield in the 400 to 450 foot interval (Unit V of the present study) of the Anchor Gas Well in Cape May. Bulliopsis integra (Conrad) and Terebra inornata Whitfield are diagnostic of the Middle Miocene St. Marys Formation, the youngest member of the Chesapeake Group.

Four transgressive sequences are interpreted from palynological assemblages on the basis of spore/dinoflagellate ratios (Fig. 3) and relative frequency distributions of continental and marine palynomorphs. The two Lower Miocene transgressions (Unit I and II) are correlated with the Tampa and Hawthorn transgressions of South Carolina, Georgia and Florida. The Middle Miocene transgressions (Units III and V) are correlated with the Calvert and St. Marys transgressions of Maryland, Delaware and Virginia. Unit IV is correlated with the Choptank Formation of Maryland, Delaware and Virginia on the basis of their stratigraphic positions and paleoenvironmental interpretations. Gernant (1970) reported that the lower portion of the Choptank Formation represents a continuation of the shoaling that began in the Calvert,

whereas the upper strata were deposited in slightly deeper water that represented the onset of the St. Marys transgression. The same paleoenvironmental interpretations are reported in this study for the sediments of Unit IV on the basis of their palynological assemblages.

Correlations of the subsurface phases of the Kirkwood Formation with other formations along the Atlantic Coastal Plain are presented in Figure 5.

AGE	N.J.	Md., Del. and Va.	N.C.	S.C.	Ga.	Fla.
UPPER MIOCENE	Cohansey Fm.	Yorktown Fm.		Duplin Marl		
				Raysor Marl		
MIDDLE MIOCENE	K I R K W O O D F O R M A T I O N	V	St. Mary's Fm.			Hawthorn Fm.
		IV	Choptank Fm.			
		III	Calvert Fm.	Pungo River Fm.		
LOWER MIOCENE		II			Hawthorn Fm.	
		I				

Figure 5: Correlation of the subsurface phases of the Kirkwood Formation with other formations along the Atlantic Coastal Plain

CORRELATION OF THE SURFACE MEMBERS AND SUBSURFACE PHASES OF
THE KIRKWOOD FORMATION

The Asbury Park and Alloway Clay Members are correlated with Unit III (Figs. 6 and 7) on the basis of their Pinus/Picea ratios, spore/dinoflagellate ratios, abundances of Gramineae and lack of Compositae and/or Abies in their palynological assemblages. The diatom, Actinoptychus heliopelta Grunow that is diagnostic of the Calvert Formation in Maryland, Delaware and Virginia is present in assemblages of Unit III as well as in those of the Asbury Park and Alloway Clay.

Surface samples of the Grenloch Sand Member are void of palynomorphs due to their coarse grain sizes and highly oxidized state. This member is correlated with Unit III of the Kirkwood Formation in the subsurface because of its interfingering stratigraphic relationship with the Asbury Park and Alloway Clay Members (Fig. 2).

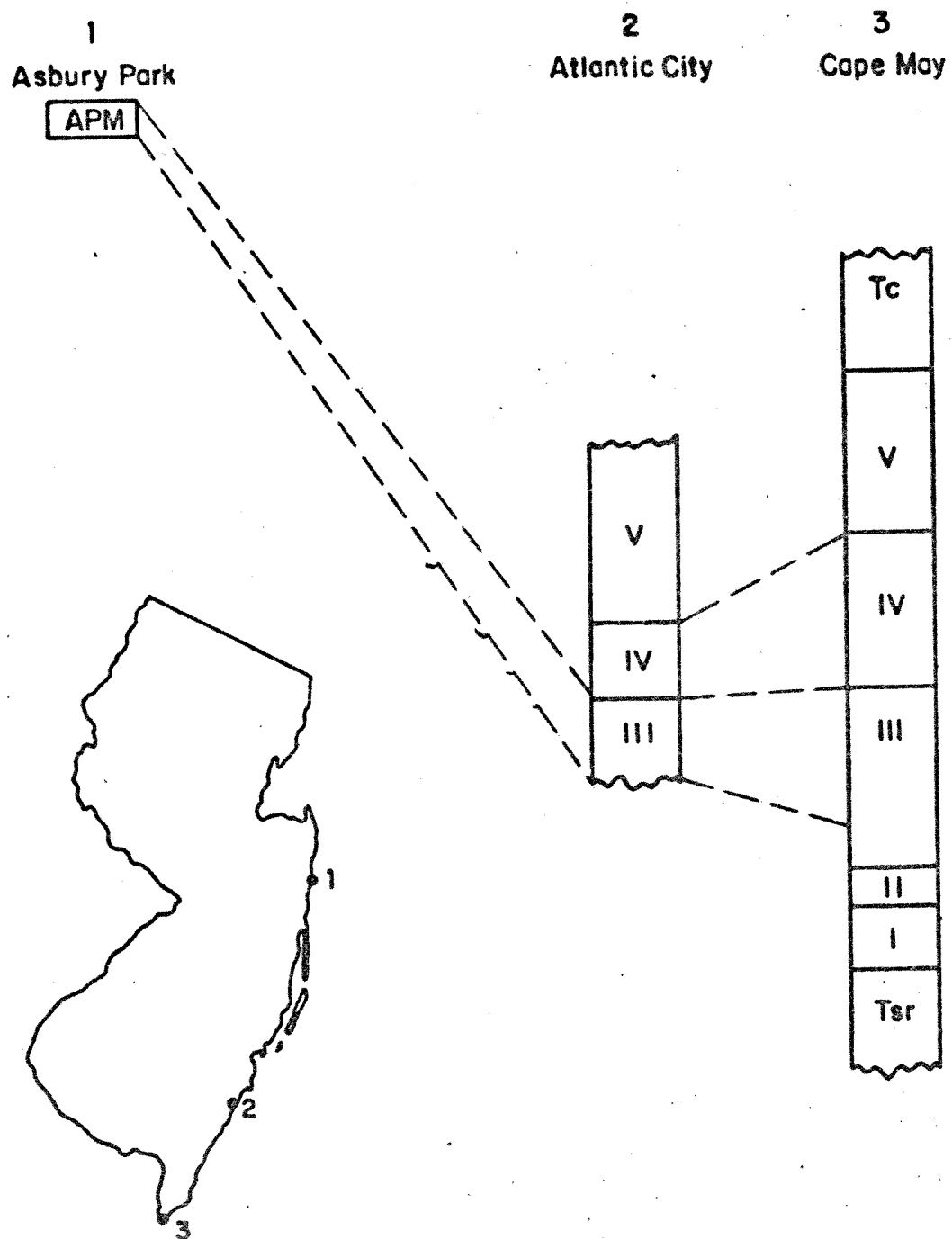


Figure 6: Correlation of the Asbury Park Member with the subsurface phases of the Kirkwood Formation

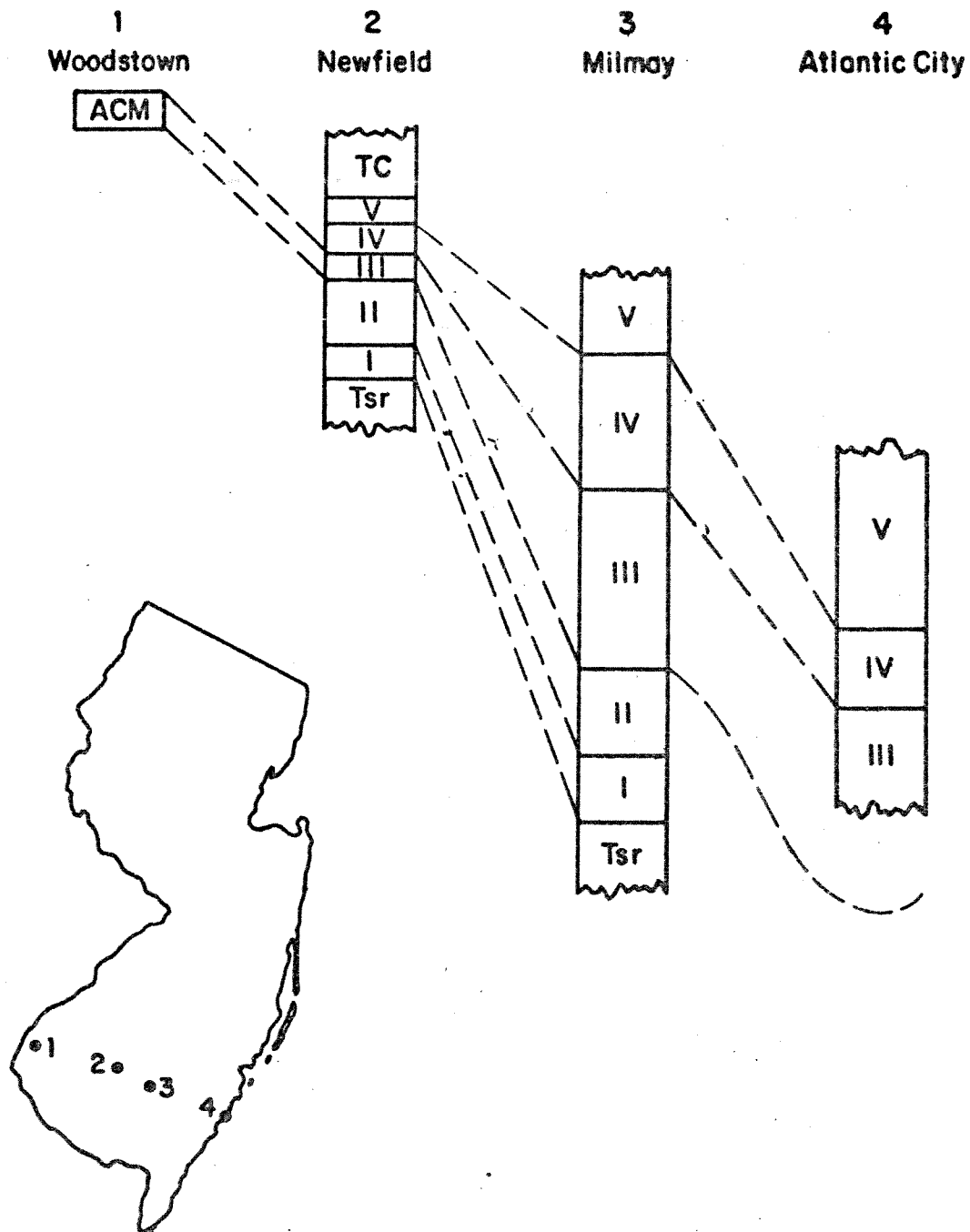


Figure 7: Correlation of the Alloway Clay Member with the subsurface phases of the Kirkwood Formation

SUMMARY OF THE SURFACE MEMBERS OF THE KIRKWOOD

MEMBER	DESCRIPTION
Asbury Park	<ol style="list-style-type: none"> 1) Calvert Equivalent 2) Deteriorating Climate (Subtropical to Temperate) 3) <u>Actinoptychus heliopelta</u> Grunow Is Present. 4) Coastal Brackish Water Environment of Deposition
Grenloch Sand	<ol style="list-style-type: none"> 1) Calvert Equivalent 2) Deteriorating Climate (Subtropical to Temperate) 3) Void of Palynological Remains 4) High Energy, Possibly Littoral, Environment of Deposition
Alloway Clay	<ol style="list-style-type: none"> 1) Calvert Equivalent 2) Deteriorating Climate (Subtropical to Temperate) 3) <u>Actinoptychus heliopelta</u> Grunow Is Present 4) <u>Pine Swamp Environment of Deposition</u>, Deepening to Middle Neritic and then Shoaling, Perhaps to Inner Neritic

SUMMARY OF THE SUBSURFACE PHASES OF THE KIRKWOOD FORMATION

ZONE	DESCRIPTION
V	1) St. Marys Equivalent 2) Deteriorating Climate (Subtropical to Temperate) 3) First Appearances of Compositae and <u>Abies</u> 4) <u>Transgressive</u> Sequence That Began In Zone IV Continues Into Zone V Followed By A Regressive Sequence
IV	1) Choptank Equivalent 2) Ameliorating Climate (Temperate to Subtropical) 3) <u>Picea</u> Becomes Rare In Upper Horizons 4) Regressive Sequence That Began In Zone III Continues Into Zone IV Followed By A <u>Transgressive</u> Sequence
III	1) Calvert Equivalent 2) Deteriorating Climate Becoming Slightly Warmer In The Later Stages (Subtropical to Temperate) 3) First Major Influxes Of <u>Picea</u> and <u>Gramineae</u> 4) <u>Transgressive</u> Sequence Followed By A Regressive Sequence
II	1) Hawthorn Equivalent 2) Deteriorating Climate (Subtropical) 3) Low Frequencies of <u>Picea</u> and <u>Gramineae</u> 4) <u>Transgressive</u> Sequence Followed By A Regressive Sequence
I	1) Tampa Equivalent 2) Deteriorating Climate (Subtropical) 3) <u>Globorotalia kugleri</u> Zone 4) <u>Transgressive</u> Sequence Followed By A Regressive Sequence

CONCLUSIONS

- 1) Subsurface sections of the Kirkwood Formation may be divided into five biostratigraphic units.
- 2) These units are correlated with the Lower Miocene Tampa and Hawthorn Formations of South Carolina, Georgia and Florida and the Middle Miocene Calvert, Choptank and St. Marys Formations of Maryland, Delaware and Virginia.
- 3) Spore/dinoflagellate ratios and relative frequencies of continental and marine palynomorphs indicate transgressive and regressive sequences in the Tampa, Hawthorn, Calvert and St. Marys Phases.
- 4) Pinus/Picea ratios and changes in the regional microfloras indicate that climatic deterioration from subtropical to temperate conditions took place during the deposition of the Tampa, Hawthorn and lower portion of the Calvert Phase followed by climatic amelioration during the deposition of the upper part of the Calvert Phase and throughout the deposition of the Choptank Phase. A second period of climatic deterioration is indicated during the deposition of the St. Marys Phase.
- 5) The Asbury Park Member is correlative with the subsurface Calvert Phase. It was deposited in a coastal brackish water environment of deposition under subtropical to temperate climatic conditions.

- 6) The Grenloch Sand Member is correlative with the subsurface Calvert Phase. It was deposited in a high energy, possibly littoral, environment of deposition under subtropical to temperate climatic conditions.
- 7) The Alloway Clay Member is correlative with the subsurface Calvert Phase. It was deposited in a low energy nearshore environment of deposition under subtropical to temperate climatic conditions.

**GEOLOGIC AGE AND CHESAPEAKE GROUP
EQUIVALENTS OF THE KIRKWOOD FORMATION, N.J.**

	1	2	3	4	5	6	7	8	9	10
INFORMAL ZONE	WOODSTOWN ALLOWAY CLAY SUBSURFACE	PITMAN ALLOWAY CLAY SUBSURFACE	GREENWICH SHILOH MARL SUBSURFACE	CAPE MAY GREENLOCH SAND SUBSURFACE	NEWFIELD GREENLOCH SAND SUBSURFACE	ATLANTIC GREENLOCH SAND SUBSURFACE	ASBURY PARK MEMBER SURFACE	ASBURY PARK MEMBER SURFACE	ASBURY PARK MEMBER SURFACE	MILMAY GREENLOCH SAND SUBSURFACE
V			C H FIRST COMPOSITAE	C H FIRST COMPOSITAE	C H FIRST COMPOSITAE	C H FIRST COMPOSITAE				C H FIRST COMPOSITAE
IV			S A P E A K	S A P E A K	S A P E A K	S A P E A K				S A P E A K
III	CHESAPEAKE FAUNA CALVERT DIATOMS	CHESAPEAKE FAUNA CALVERT DIATOMS	CALVERT SILICO-FLAGELLATE	F A U N A	F A U N A		CALVERT DIATOMS	CALVERT DIATOMS	CALVERT DIATOMS	F A U N A
II					GLOBE - ROTALLIA KUGLERI ZONE BASAL A MIOCENE					F A U N A
I										F A U N A
			EOCENE							EOCENE

Figure 8: Geologic ages and Chesapeake Group equivalents of the surface and subsurface facies of the Kirkwood Formation

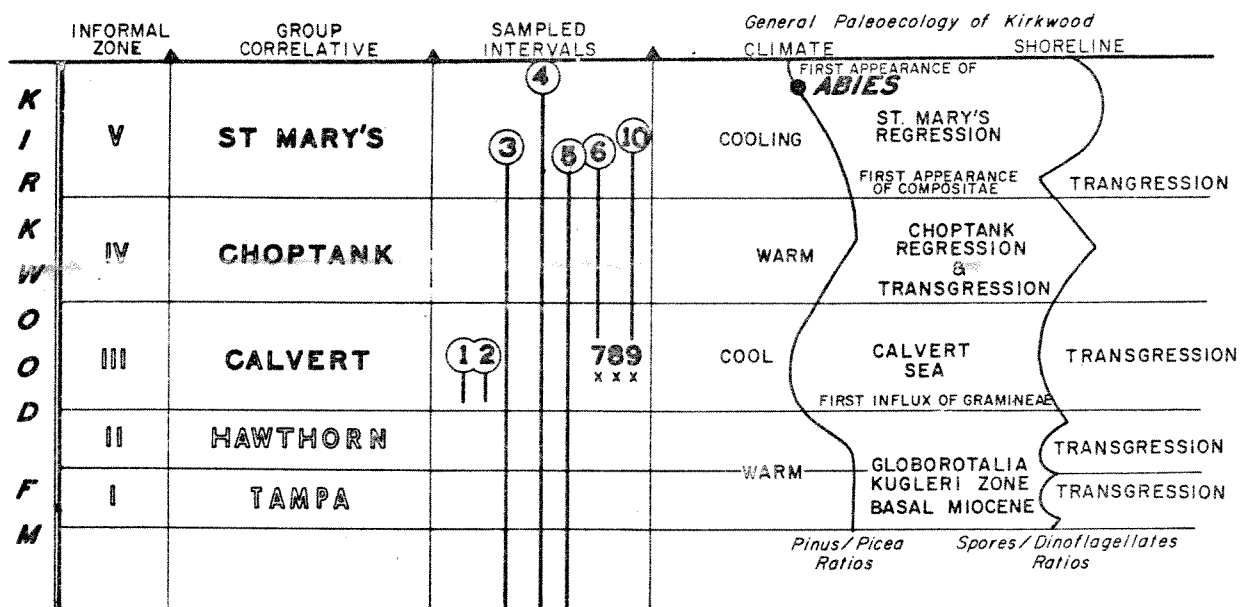


Figure 9: Correlations and paleoenvironmental inferences of the surface and subsurface facies of the Kirkwood Formation

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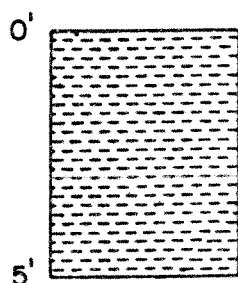
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APPENDIX I

Surface and Subsurface Sections of the Kirkwood Formation

Farmingdale Sample

1 1/4 miles north of Farmingdale, about 200 yards east of the railroad tracks, at an elevation of approximately 15 feet above sea level



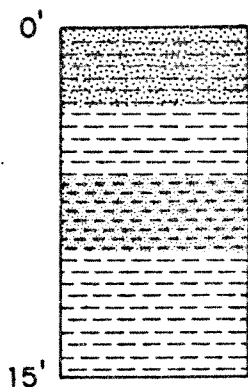
(Unit III)

Black Clay

Chenopods, Equisetum, Ericaceae, Osmunda, and Polypodium dominate assemblages

Hammonton Sample

south side of Hammonton Road, about 2 miles east of Route 9, at an elevation of approximately 25 feet above sea level



(Unit III)

Light Gray Sands and Clays--Low Frequencies of Palynomorphs

Light Gray Clay--Low Frequencies of Palynomorphs

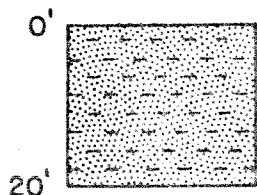
Dark Black Silty Clay--Nonarboreal elements dominate assemblages

Light Gray Clay--Low Frequencies of Palynomorphs

Light Gray Clay--Low Frequencies of Palynomorphs

Howell Sample

landfill site, about 3 miles east of Route 9 and 1 mile north of the Ocean County Line, at an elevation of approximately 40 feet above sea level



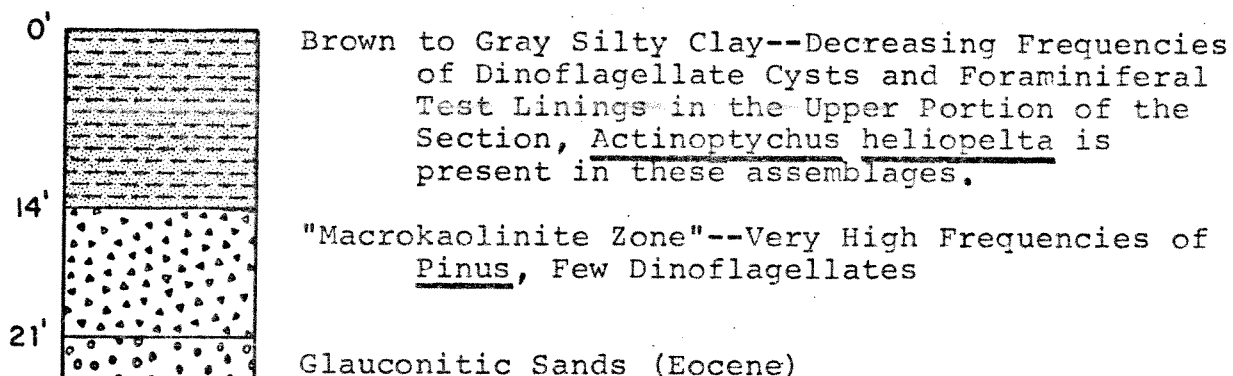
(Unit III)

Red Sands with small Gray Clay Inclusions--Low Frequencies of Palynomorphs and Actinopterychus heliopelta

Figure 10: Surface Samples of the Asbury Park Member

Woodstown Sample

(Unit III)



Pitman Sample

(Unit III)

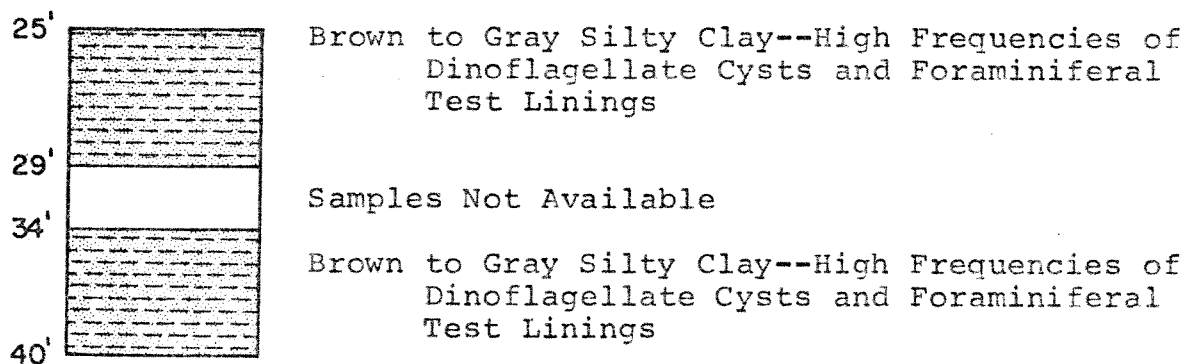


Figure 11: Subsurface Samples of the Alloway Clay Member

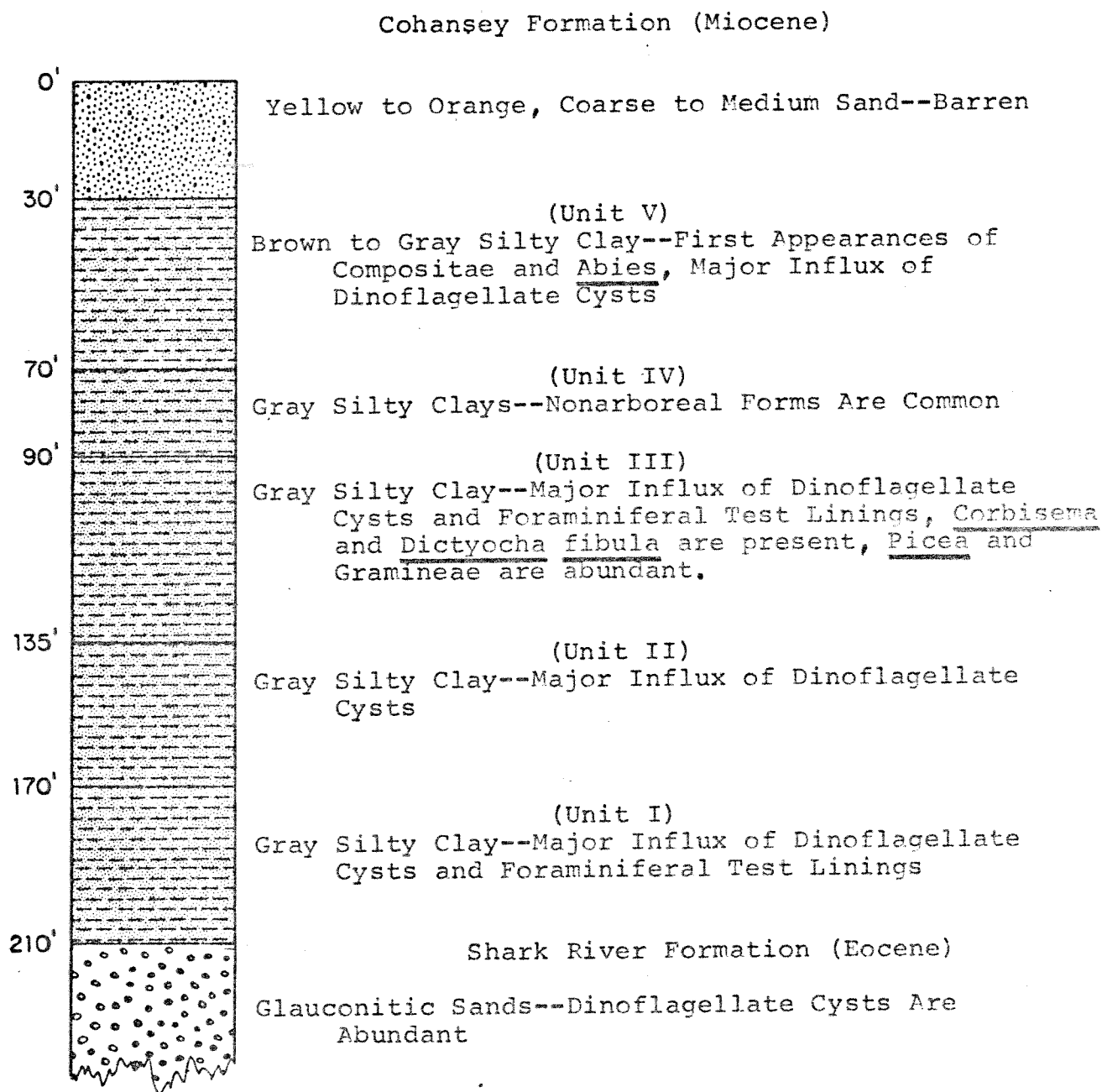


Figure 12: Subsurface Section of the Kirkwood Formation From Greenwich, New Jersey

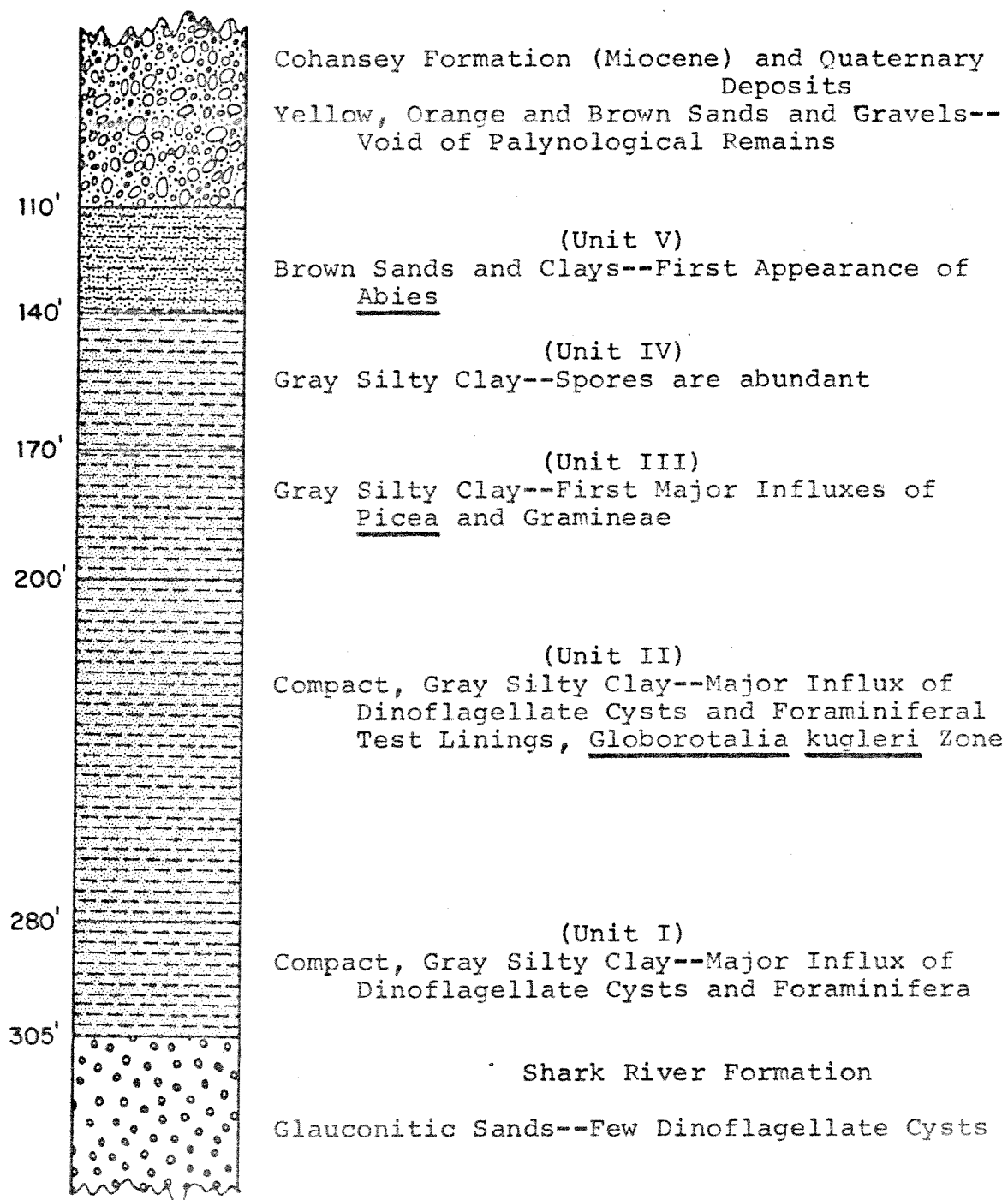


Figure 13: Subsurface Section of the Kirkwood Formation
from Newfield, New Jersey

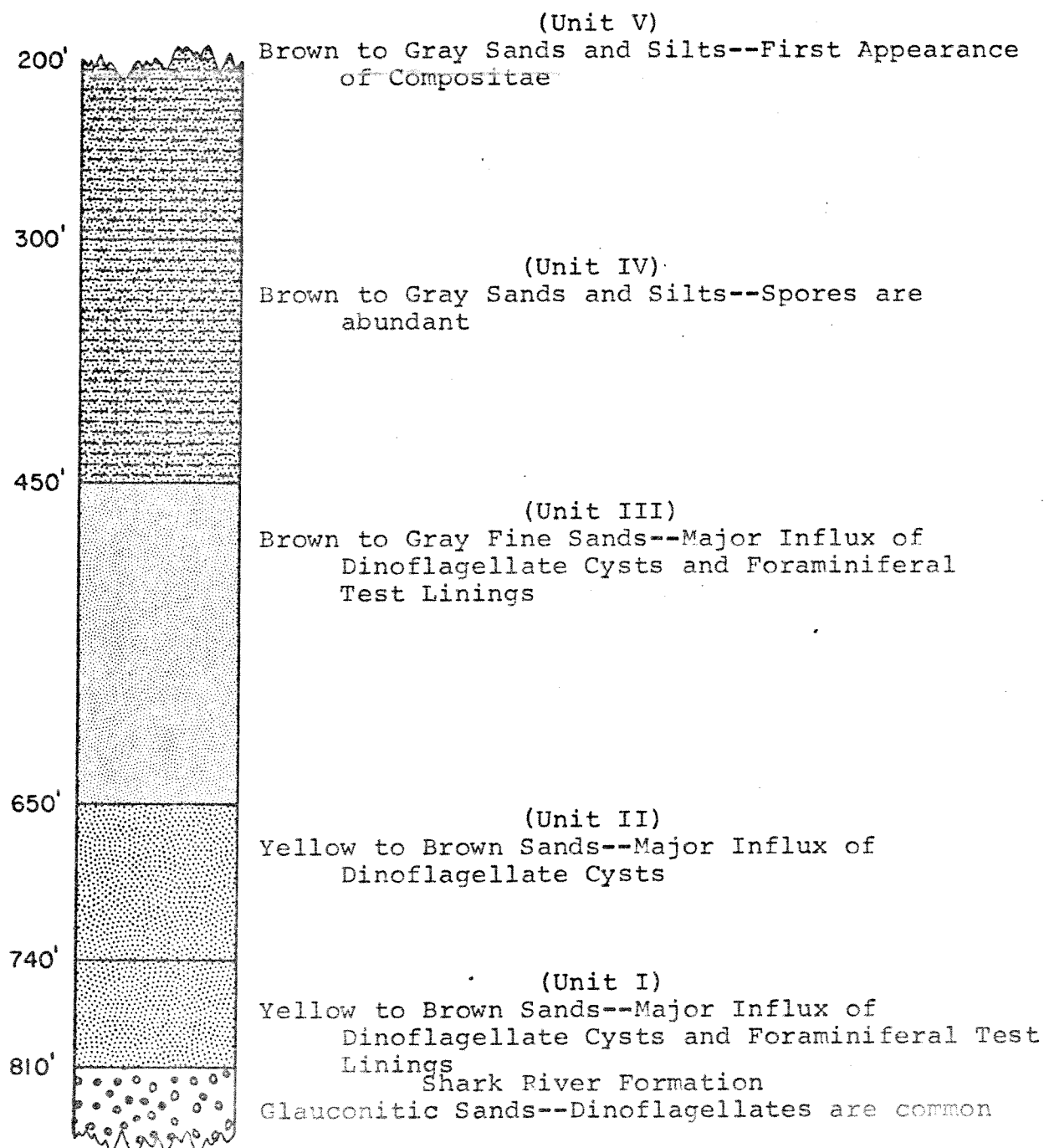


Figure 14: Subsurface Section of the Kirkwood Formation from Milmay, New Jersey

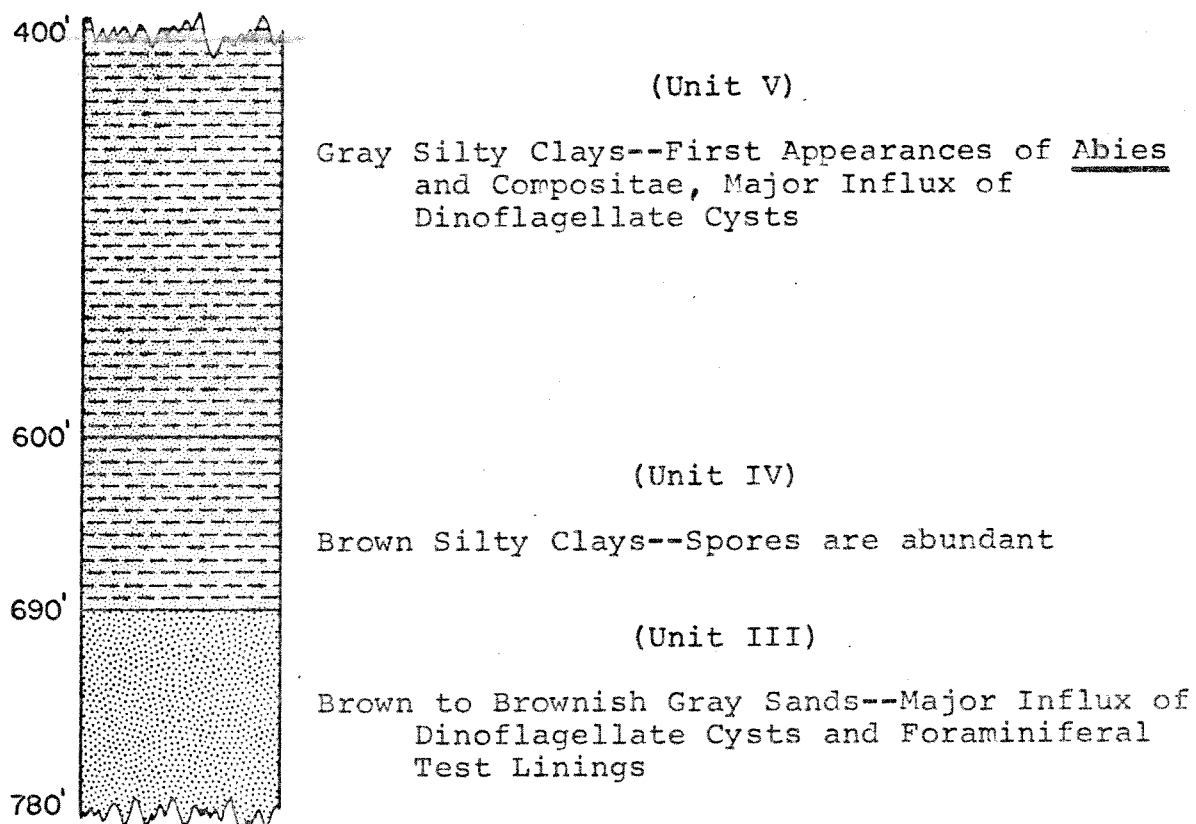


Figure 15: Subsurface Section of the Kirkwood Formation from Atlantic City, New Jersey

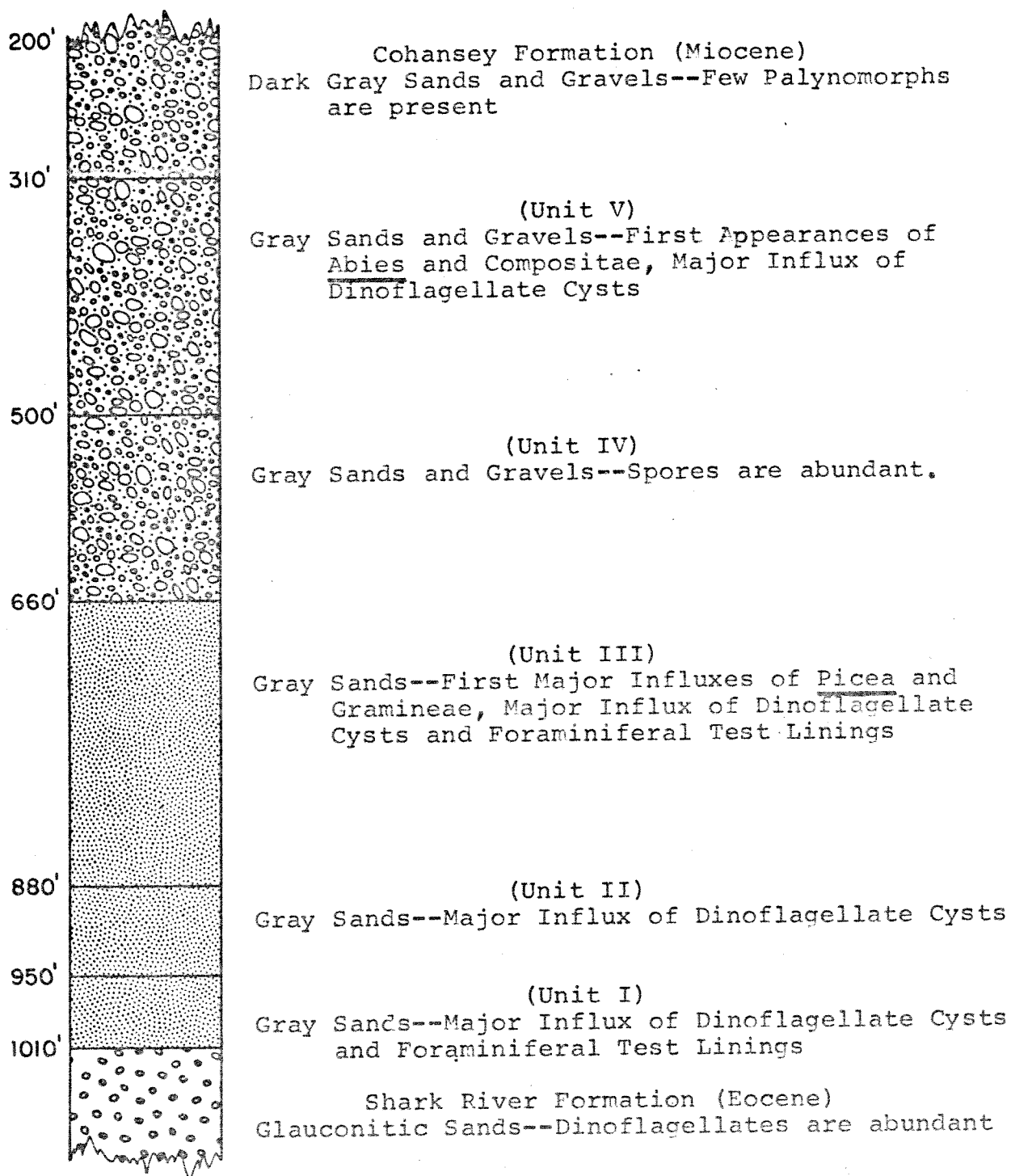
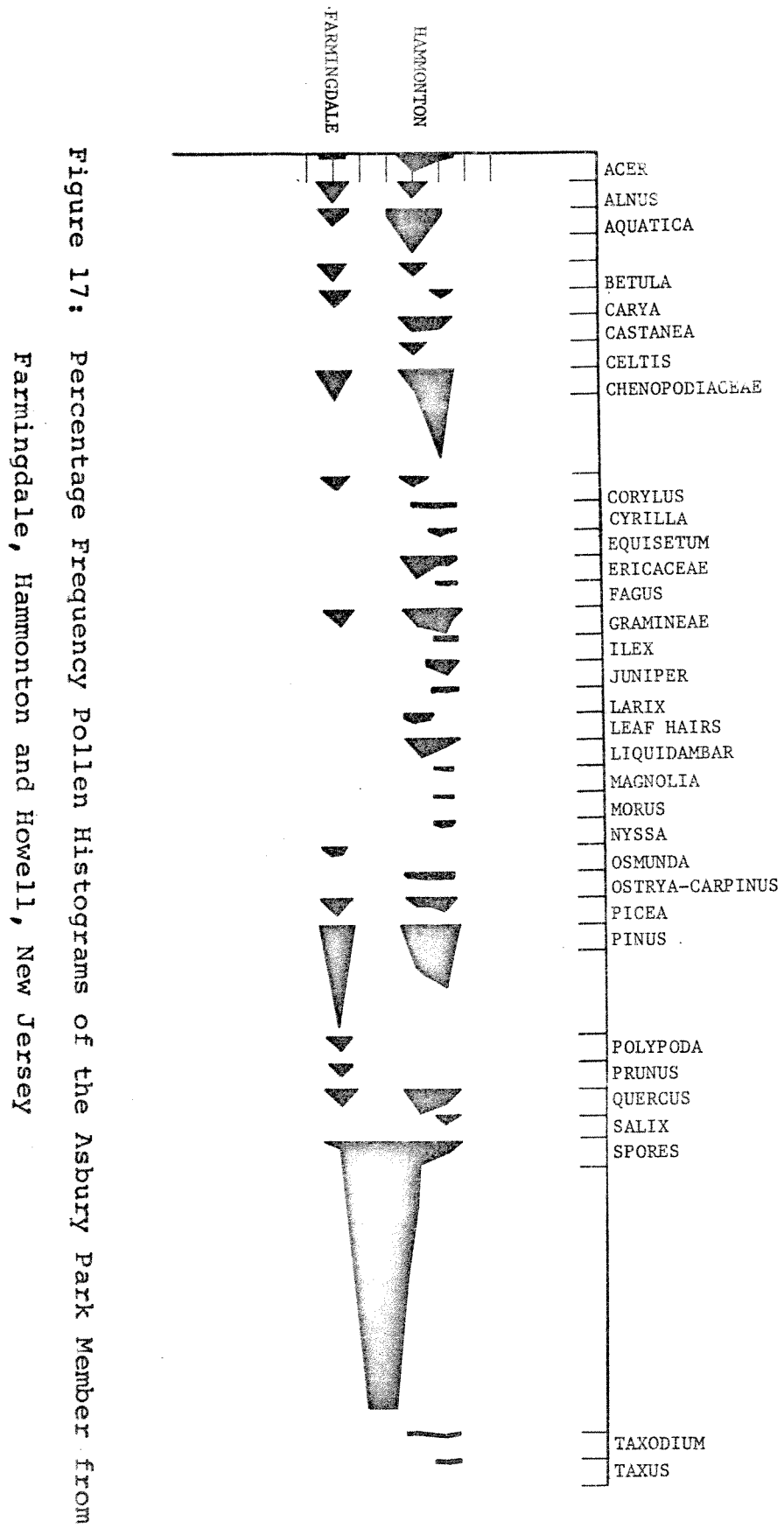


Figure 16: Subsurface Section of the Kirkwood Formation from Cape May, New Jersey

APPENDIX II

Percentage Frequency Pollen Histograms of the Kirkwood Formation



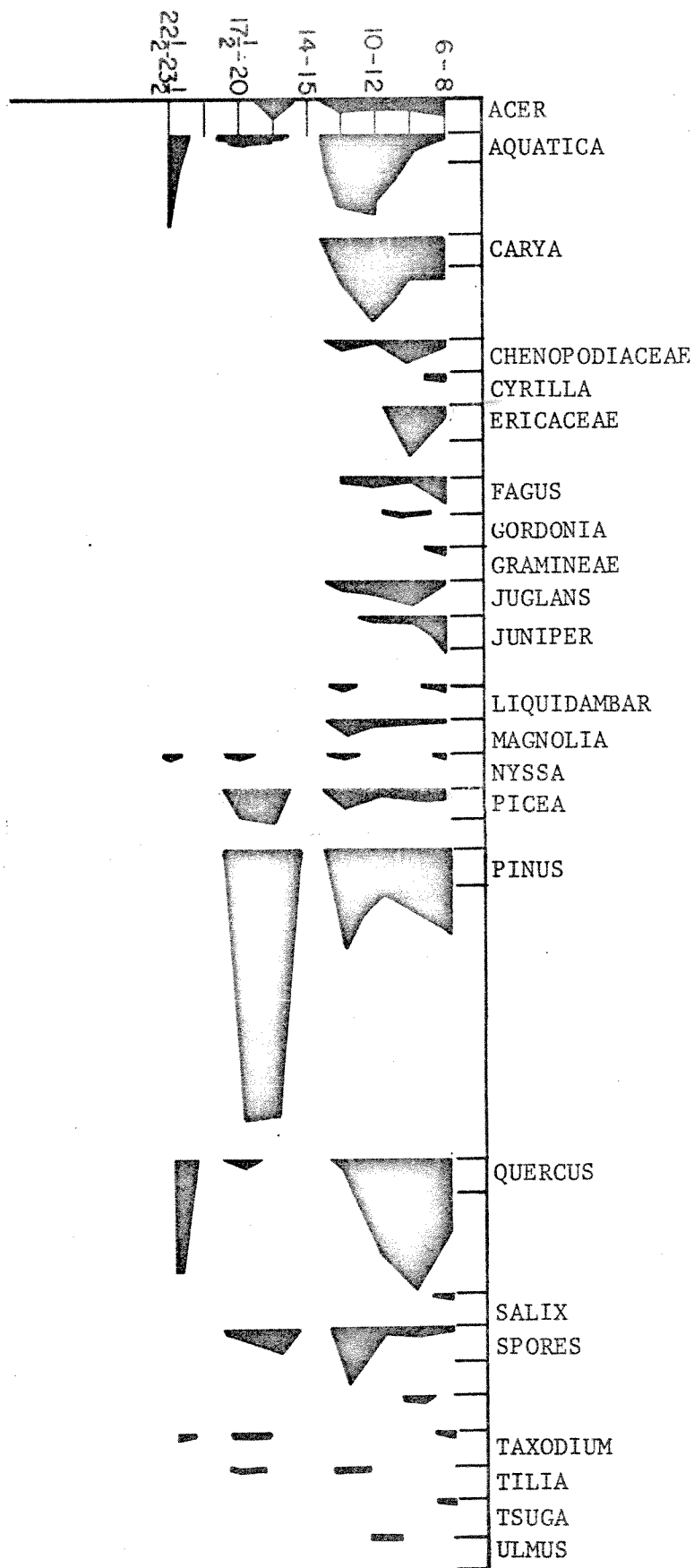


Figure 18: Percentage frequency pollen histogram of the Alloway Clay Member from
Woodstown, New Jersey

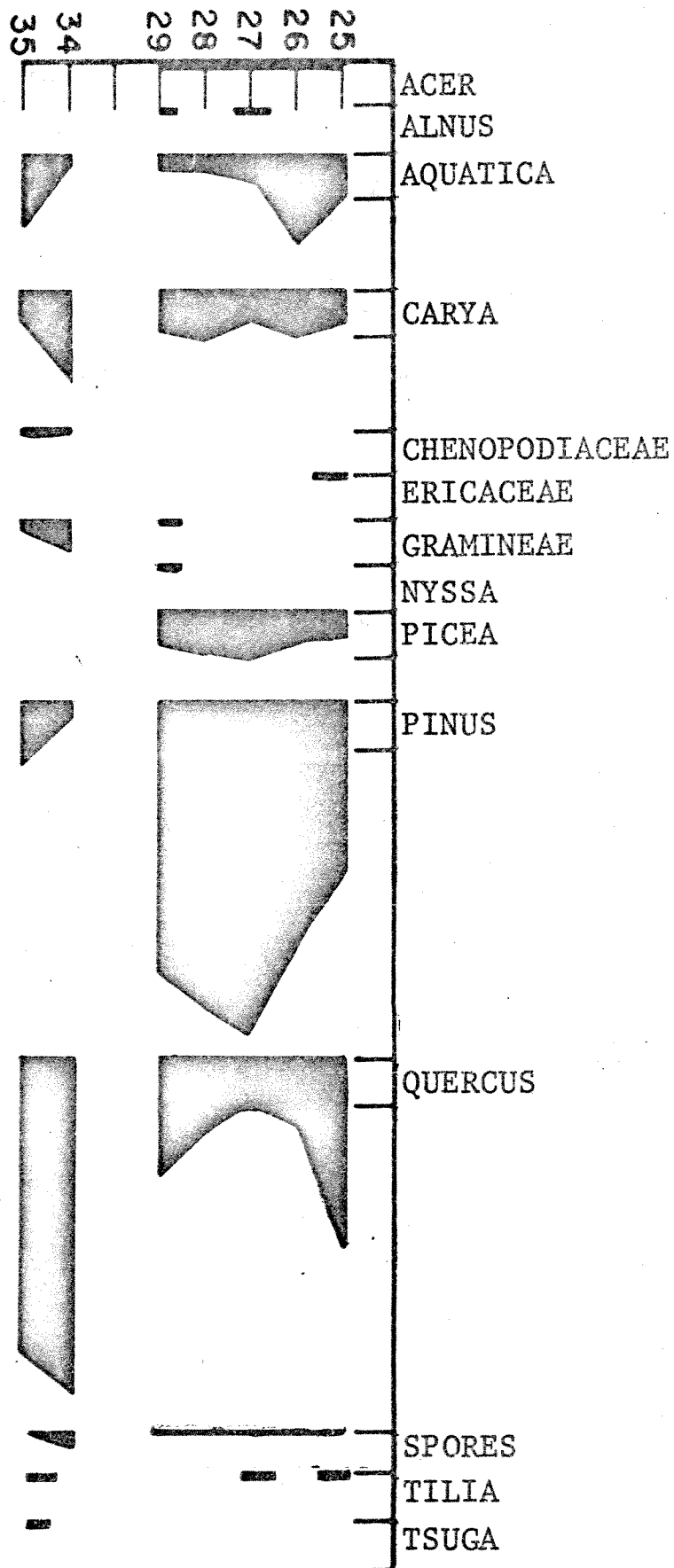
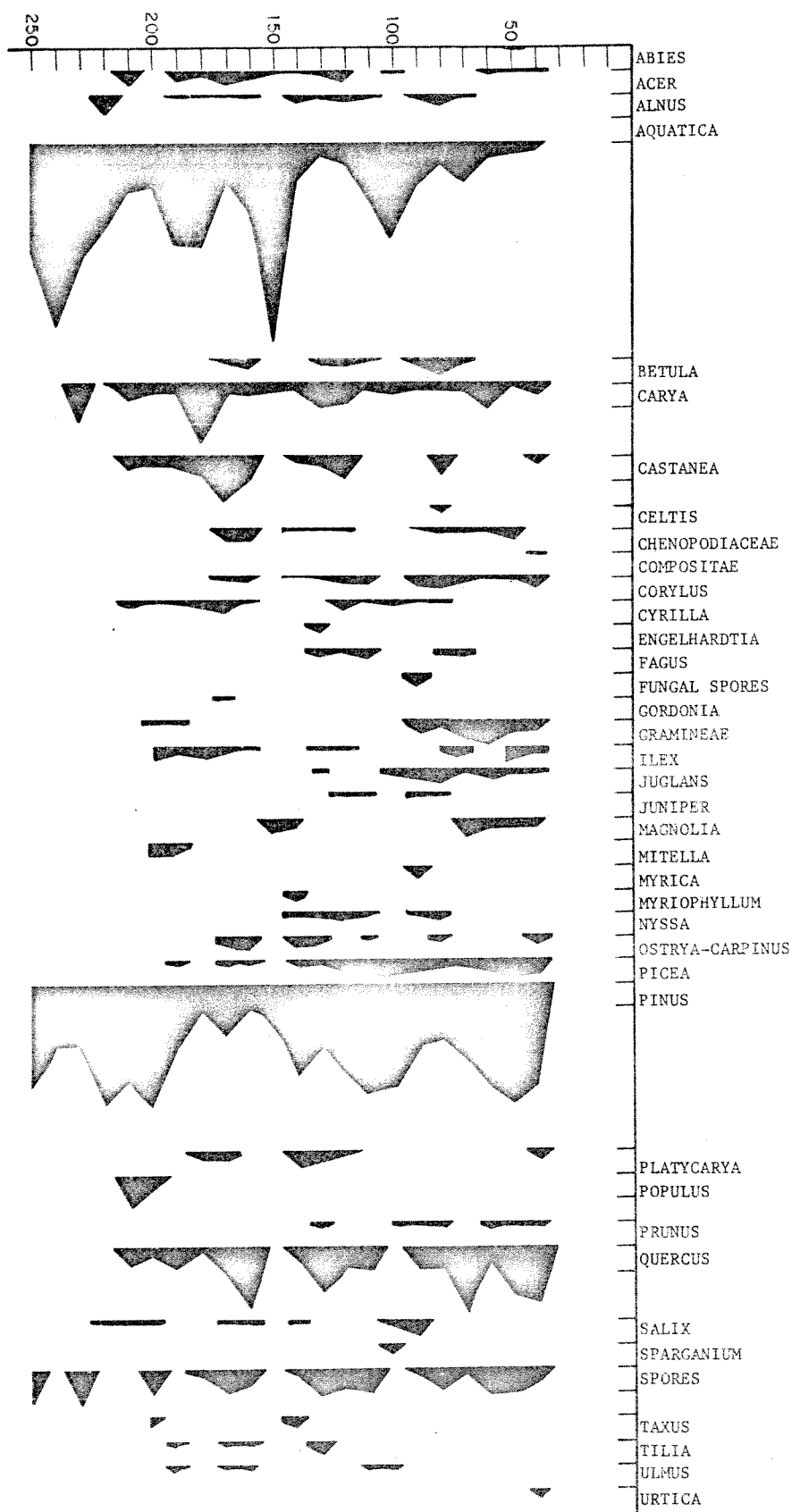


Figure 19: Percentage frequency pollen histogram of the Alloway Clay Member from Pitman, New Jersey

Figure 20: Percentage frequency pollen histogram of the Kirkwood Formation from Greenwich, New Jersey



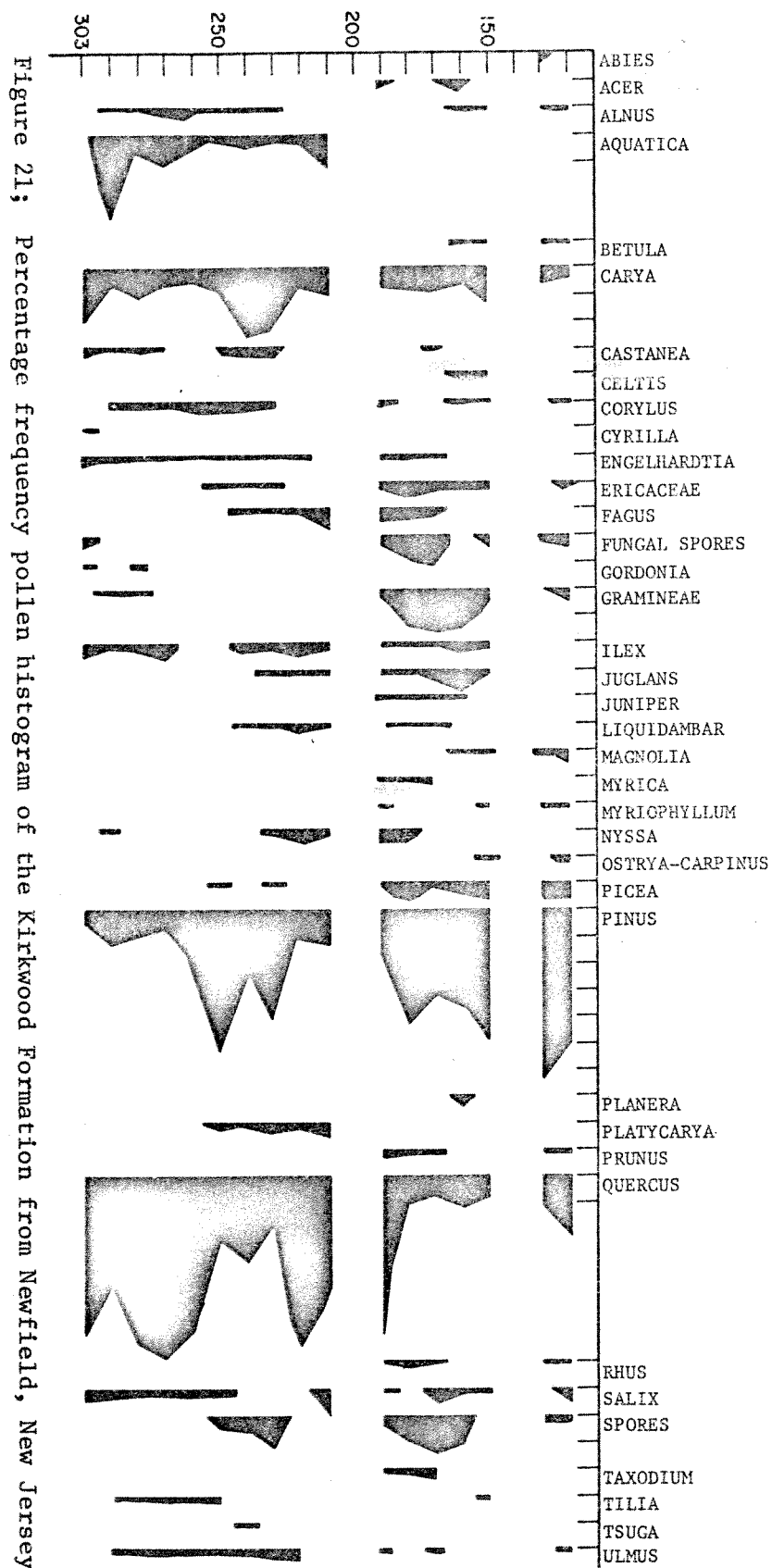


Figure 21; Percentage frequency pollen histogram of the Kirkwood Formation from Newfield, New Jersey

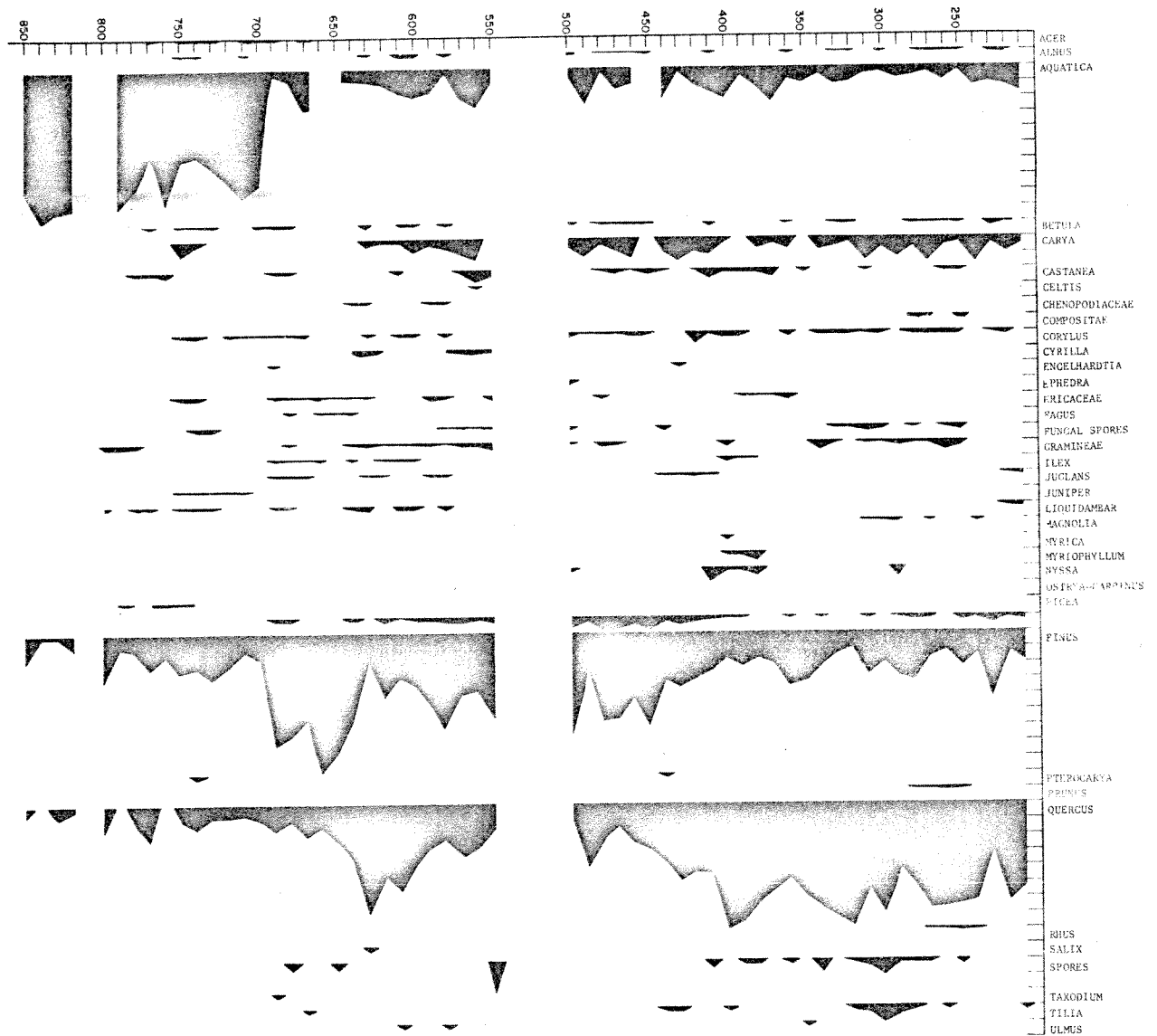


Figure 22: Percentage Frequency Pollen Histograms of the Kirkwood Formation from Milmay, New Jersey

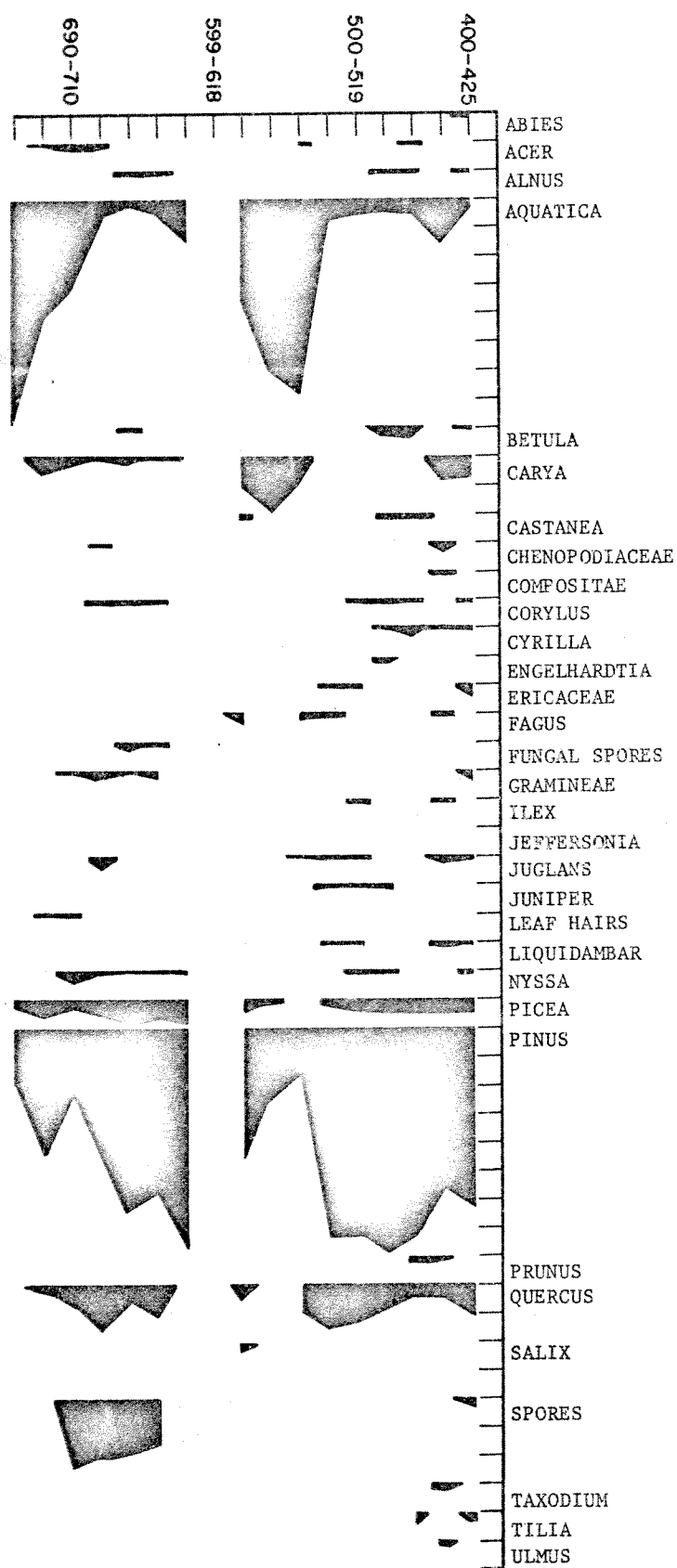
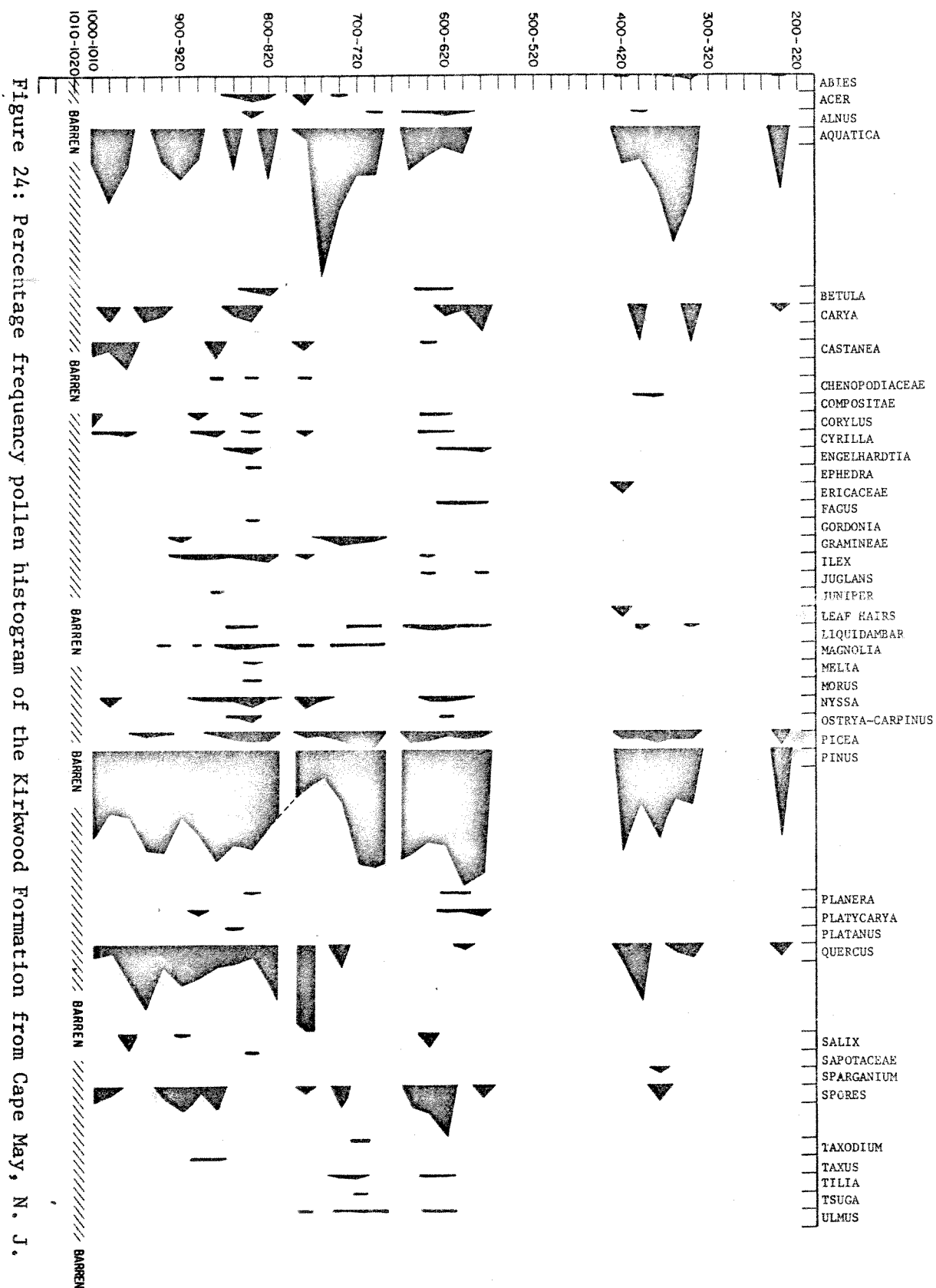


Figure 23: Percentage frequency pollen histogram of the Kirkwood Formation from
Atlantic City, New Jersey



APPENDIX III

Absolute Frequency Counts of Palynomorphs and other Microfossil
Groups of the Kirkwood Formation

WOODSTOWN, NEW JERSEY

[illegible]

Figure 26: Absolute Frequency Count of Palynomorphs and other Microfossil Groups of the Alloway Clay Member from Woodstown, New Jersey

NEWFIELD, NEW JERSEY

[illegible]

Figure 29: Absolute Frequency Count of Palynomorphs and other Microfossil Groups of the Kirkwood Formation from Newfield, New Jersey

Figure 30: Absolute Frequency Count of Palynomorphs and other Microfossil Groups of the Kirkwood Formation from Milway, New Jersey

MILWAY, NEW JERSEY		PALYNOFORMS AND OTHER MICROFOSSIL GROUPS										TOTAL		PERCENTAGE	
		ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	ALVEOLATE	TOTAL	PERCENTAGE	PERCENTAGE	PERCENTAGE
200	200	1	1	1	1	1	1	1	1	1	1	10	10	10	10
250	250	1	1	1	1	1	1	1	1	1	1	10	10	10	10
300	300	1	1	1	1	1	1	1	1	1	1	10	10	10	10
350	350	1	1	1	1	1	1	1	1	1	1	10	10	10	10
400	400	1	1	1	1	1	1	1	1	1	1	10	10	10	10
450	450	1	1	1	1	1	1	1	1	1	1	10	10	10	10
500	500	1	1	1	1	1	1	1	1	1	1	10	10	10	10
550	550	1	1	1	1	1	1	1	1	1	1	10	10	10	10
600	600	1	1	1	1	1	1	1	1	1	1	10	10	10	10
650	650	1	1	1	1	1	1	1	1	1	1	10	10	10	10
700	700	1	1	1	1	1	1	1	1	1	1	10	10	10	10
750	750	1	1	1	1	1	1	1	1	1	1	10	10	10	10
800	800	1	1	1	1	1	1	1	1	1	1	10	10	10	10

Fredric Robert Goldstein

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- 1966 B.S. in Geology, Brooklyn, New York
- 1966-68 Attended Miami University, Oxford, Ohio
- 1966-68 Teaching Assistantship, Miami University
- 1967 June-August, Instructor, Department of Geology,
Brooklyn College
- 1968 Publication: A Geological Guide of the Bronx Park,
New York Zoological Publication
- 1968 M.S. in Geology, Miami University
- 1968-69 Attended Cornell University, Ithaca, New York
- 1968-69 Teaching Fellowship, Cornell University
- 1969-74 Graduate Work in Geology, Rutgers University,
New Brunswick, New Jersey
- 1969-71 Teaching Assistantship, Rutgers University
- 1970 Abstract: Initial Results of the Palynology of
the Kirkwood Formation, Presented at the Annual
Meeting of the New Jersey Academy of Science,
Princeton University, Princeton, New Jersey
- 1971-72 Instructor, Science Department, Lakewood High
School, Lakewood, New Jersey
- 1972-73 Instructor, Department of Earth, Space and
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- 1973 Abstract: The Palynology of the Kirkwood
Formation, Presented at the Northeastern Sectional
Meeting of the Geological Society of America,
Allentown, Pennsylvania
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- 1973- Assistant Professor, Department of Chemistry,
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